

## Introduction to Salt Dilution Gauging for Streamflow Measurement Part IV:

# The Mass Balance (or Dry Injection) Method

Rob Hudson and John Fraser

In part one of this series, Moore (2004a) introduced the general principles of stream gauging by salt dilution. In subsequent articles, Moore (2004b, 2005) described techniques of constant-rate injection and slug injection using salt in solution. This is the final article in the series and details the “mass balance method.” Originally described by Elder et al. (1990), the mass balance method differs from slug injection using salt in solution (Moore 2005) in that it is based on conservation of tracer mass, not of tracer volume. By using the mass balance method, salt can be injected into the stream either in dry form or in solution. Because it is more common to work with dry salt, this method has also become known as the “dry injection method.”

Dilution methods have been used for at least five decades (Østrem 1964; Church and Kellerhals 1970; Church 1975; Day 1976, 1977a, 1997b; Hongve 1987; Johnstone 1988; Kite 1993). In 1998 we began to develop the mass balance method for use in small BC streams where current metering is

either difficult or impossible (Hudson and Fraser 2002). Subsequently, we have applied this method successfully in several coastal streams including Russell Creek (Hudson and Fraser 2002), upper Nahmint River on Vancouver Island, Flume Creek (Sunshine Coast), and Culliton and Furry Creeks (Sea-to-Sky Highway).

### Background

The basic principle of dilution gauging is to add a known quantity of a tracer to a stream and observe its concentration in the stream at a point where it is fully mixed with the flow. The higher the flow, the more it dilutes the tracer. Dry salt used as the tracer must be injected at a point that favours rapid dissolution. This creates a salt solution *in situ* that then disperses into the flow aided by turbulence in the water column. The resulting concentration of salt is measured as electrical conductivity at a point downstream of the injection point where it is completely mixed. The distance between the injection and measurement points is known as the mixing length ( $L$ ). The dispersion pattern of conductivity over time is similar in shape to a storm hydrograph (Figure 1).

*The mass balance method is based on conservation of tracer mass, not of tracer volume.*

Streamflow  $Q$  is calculated using equation 1 where  $M$  is the mass of salt (in grams) and  $A$  is the area under the graph of concentration over time (Figure 1). The units of  $A$  are milligram-seconds per litre (equivalent to  $g \cdot s/m^3$ ).

$$Q = \frac{M}{A} \quad (1)$$

The quantity  $A$  in equation 1 and shown on the graph (Figure 1) can be calculated as:

$$A = \int c_t \, dt \quad (2)$$

where  $c_t$  is the concentration of injected salt at time  $t$ , and  $t_{int}$  is the time interval between successive data points.

As noted above, the salt concentration is measured as electrical conductivity ( $EC$ ) in the stream. The concentration of the injected salt can be calculated using equation 3 below:

$$c_t = (EC_t - EC_0) \cdot CF \quad (3)$$

where  $EC_t$  is the electrical conductivity at time  $t$ ,  $EC_0$  is the baseline conductivity, and  $CF$  is the concentration factor. The concentration factor is the coefficient in the near-linear relation between  $EC$  and salt concentration. However,  $CF$  is not a constant, since stream temperature and background chemistry also affect  $EC$ . These effects constitute a potential source of error that can only be controlled by understanding the relationships between  $EC$  and salt concentration, stream temperature, and chemistry.

### Factors Affecting $EC$ versus Salt Concentration<sup>1</sup>

For a salt dilution measurement, both the background ion concentration in the stream and the concentration of salt added to the stream affect the  $CF$ . Hongve (1987) found that the

<sup>1</sup> Within the salt dilution gauging literature, there is a lack of agreement about the derivation of a unified calibration approach. We have developed the procedures presented in this article as one example of an approach that could be used and (or) modified by the reader. These procedures are based on trials that we conducted for equipment, streamflows, and watersheds we commonly sampled. As with any measurement technique, procedural steps and calibration assumptions should be validated for the local watersheds in which the technique is applied.

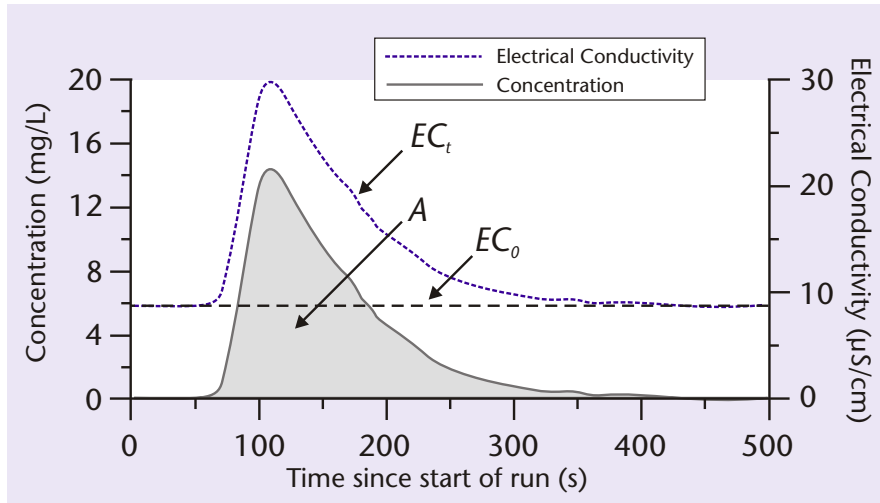


Figure 1. Salt dilution measurement at Russell Creek, October 2003. The shaded area is the quantity A that must be calculated.

concentration factor was directly related to the baseline conductivity of the stream. For a pure NaCl solution, 1 mg of NaCl added to 1 L of de-ionized water will increase conductivity by 2.14  $\mu\text{S}/\text{cm}$  ( $\text{CF} = 0.467$ ). As the ionic strength of the solution increases, the interaction of ions in solution begins to hinder each others' activity. The CF is equal to 0.47 for [NaCl] in the range from 0 to 30 mg/L (EC from 0 to 64  $\mu\text{S}/\text{cm}$ ) and increases to 0.51 for [NaCl] in the range from 300 to 1000 mg/L (EC from 617 to 1990  $\mu\text{S}/\text{cm}$ ). However, the presence of ions other than NaCl in the stream being measured may result in a slightly different range of values for the CF.

The relationship between EC and temperature is more-or-less linear in the range of temperatures commonly encountered during flow measurements, but there is considerable lack of agreement in the literature concerning that relationship at low temperatures (i.e., 0–3°C). Smart (1992) found a linear relation for temperatures ranging from less than 1 to 10°C, contradicting statements by Østrem (1964) and Collins (1978). Johnstone (1988) reported linear relations for temperatures from 0.5 to 25°C.

We have spent considerable effort in trying to calibrate our conductivity probes and believed we had a reasonable temperature correction curve based on a combination of field and laboratory calibrations (Figure 2). However, it has now become evident that differences in water chemistry among different streams have a significant effect on the value of CF such that partial calibrations from different sites cannot be combined. Rather, a separate temperature correction curve must be developed for each stream, and laboratory calibrations are not necessarily applicable in the field. We have therefore used a simple calibration procedure to derive a specific value of CF for each measurement. A simple calibration kit and procedure are described below.

## Probe Calibration<sup>2</sup>

### Calibration kit

The calibration kit should consist of the following items:

1. A 1-L graduated cylinder
2. A 10-mL glass pipette and filler bulb
3. Standard solutions of 2000 and 20 000 mg/L NaCl
4. Distilled or de-ionized water

### Calibration procedure

This procedure should be performed each time a measurement is collected.

1. Allow the EC probe to equilibrate to stream temperature. Note the EC and the temperature.
2. Obtain a 980-mL sample of the streamwater in the graduated cylinder.
3. Place the probe in the cylinder and note the EC and temperature. These values should be close or identical to the values obtained in step 1. If so, they represent baseline values. If not, get a different sample.

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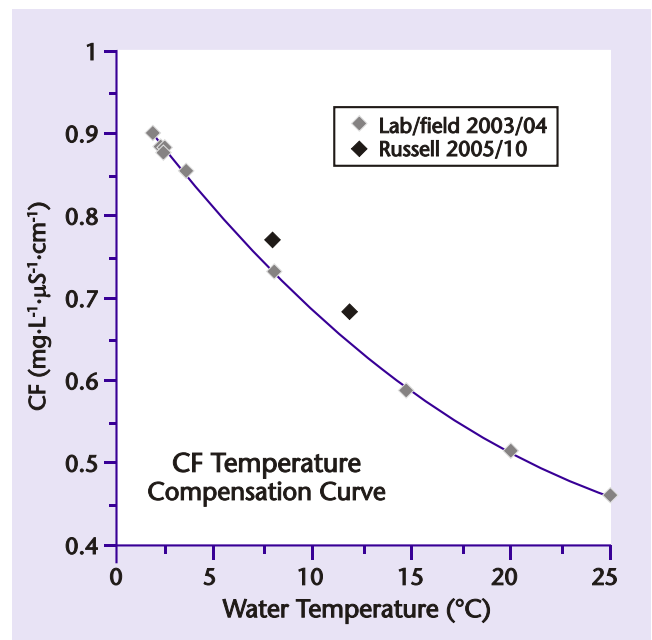


Figure 2. This CF temperature curve was developed in the lab and at a field site (Furry Creek) where baseline EC was very low (about 10  $\mu\text{S}/\text{cm}$ ). However, as noted by Hongve (1987), the stream chemistry affects the value of CF such that at Russell Creek, where baseline EC was higher, the CF is also slightly higher.

<sup>2</sup> The procedures presented illustrate one approach to probe calibration; other approaches may be used or need to be developed in the application of the mass balance method to meet unique watershed conditions.

4. Add 20 mL of standard solution to the cylinder in two 10-mL increments. Note the EC and temperature after each increment. The choice of standard depends on the expected range of the peak concentration measured in the stream. Concentrations at each step are the concentration above background.

- a. Range 1 = 0–40 mg/L using standard 1 at 2000 mg/L NaCl
  - i. Step 1 concentration = 20.2 mg/L
  - ii. Step 2 concentration = 40.0 mg/L
- b. Range 2 = 0–400 mg/L using standard 2 at 20 000 mg/L NaCl
  - i. Step 1 concentration = 20.2 mg/L
  - ii. Step 2 concentration = 40.0 mg/L

5. Calculate the CF for each step.

$$CF = \Delta(\text{Concentration}) / (EC - EC_{\text{base}})$$

These readings should be tabulated and maintained over the long term. This will result in a record of temperature-specific CF measurements that can be used to derive a site-specific “CF curve.”

## Equipment

The main pieces of equipment needed to measure streamflow by salt dilution are the electrical conductivity (EC) logger, scales for weighing salt, and a calibration kit for performing simple calibrations *in situ*. A notebook computer may be needed depending on the capability of the EC logger.

### EC logger

We have used two different 4-electrode temperature-compensated EC probes with dedicated micro-loggers to collect our instream conductivity data. These instruments auto-range between 0 and 200 000  $\mu\text{S}/\text{cm}$ , with an accuracy of  $\pm 0.5\%$  of the measured value. The conductivity logger should be capable of measuring and recording stream

temperature as well as EC (primarily from 0 to 1000  $\mu\text{S}/\text{cm}$ ) at a log interval of 5 s or less. These instruments also have linear temperature compensation and data can be recorded in either compensated or raw (i.e., not corrected for temperature) form. However, we have found the temperature correction either unreliable or not useful since, as stated above, stream chemistry also affects the value of CF. Therefore we find that the raw EC data should be used and a specific value for CF should be derived through *in situ* calibration.

### Scales

The accuracy of streamflow measurement is usually expressed as a percentage of the true streamflow. In this method, the accuracy to which the injection mass of salt can be measured is clearly an important factor. At low flow (0–2  $\text{m}^3/\text{s}$ ) using 1 kg of salt for 1% accuracy, the salt mass should be known to within 10 g. A good-quality scale for weighing food is probably sufficient for this level of accuracy. For low flows requiring less than 2 kg of salt, injection masses can be pre-weighed and kept in zip-lock freezer bags. The bags should be weighed before and after use to determine the actual mass of salt used in a measurement. For high flows (> 20  $\text{m}^3/\text{s}$ ), salt can be added in multiples of 20 kg using the bags directly from the bulk distributor. The bags of salt should be weighed to within 200 g for 1% accuracy. Typically this level of accuracy can be achieved with a good-quality bathroom scale. However, for weighing salt masses suitable for probe calibration, a laboratory scale with a range of 0 to 2000 g capable of 2 digits of precision (i.e.,  $\pm 0.01$  g) is required.

## Application of Dry Injection at a Stream-gauging Site

A stream-gauging site that is not suited to current metering or other methods (weir, flume, etc.) might be

gauged by dry injection salt dilution. The following list outlines the criteria that should be considered before applying the method.

### Preliminary criteria

Evaluate the site for suitability. The basic characteristics of a reach suitable for salt dilution are:

1. Turbulent at all flows.
2. Steep gradient: Some channels with gradients between 3 and 5% can be measured with salt dilution. Low gradient (<3%) reaches tend to be suitable for current metering and high gradient (>5%) for salt dilution.
3. Minimal pools and other backwater areas.
4. No tributary inflows in the gauging reach.
5. Riffle–pool, step–pool, cascade–pool morphology with cobble–boulder bed and flow constrictions.

The above criteria are easy to evaluate at a field site. However, the most critical considerations in applying the method are:

6. Ability to perform a clean injection at a point that favours mixing and rapid dissolution.
7. The salt must be fully mixed with the flow at the point where EC is measured in the channel.

To meet criterion #6 is simply a matter of technique. For “clean injection,” all the salt is injected into a point of turbulence with a single movement. The method may still work if the salt is injected in stages, but it will be difficult to determine from the dispersion graph if the measurement has been successful. The ideal injection point is a constriction in the channel where most of the flow converges and passes between boulders in the channel bed. For example, at Stephanie Creek a series of constrictions as the channel passes under the bridge (Figure 3) makes it an ideal injection site for two reasons:





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Figure 3. At Stephanie Creek the flow passes through a series of three constrictions that mix the salt very efficiently, resulting in a short mixing length.

- the bridge can be used to dump the salt directly into the injection point; and
- the turbulence created below the constriction helps to dissolve the salt and mix the resulting solution into the water column.

Criterion #7 is more difficult to verify because it cannot be determined directly without deploying numerous probes in the channel. Choosing an appropriate mixing length is by far the most difficult aspect of the field procedure and requires an understanding of the dissolution and dispersion processes of salt in flowing water.

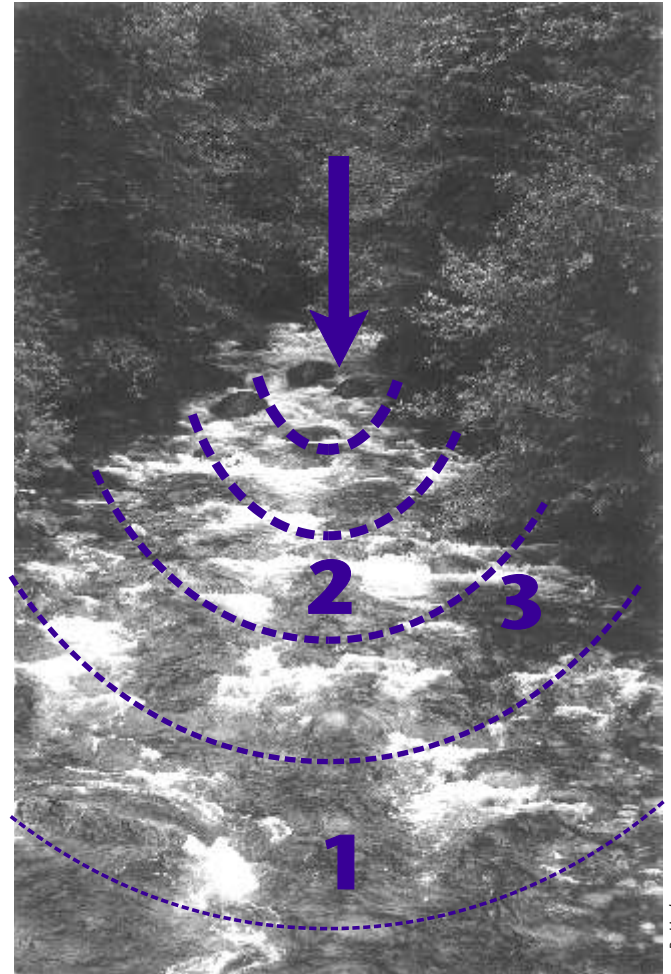
#### Behaviour of injected salt in a stream channel

The dissolution of salt in water takes time, with the rate of dissolution being proportional to water temperature (i.e., it dissolves faster in

warm than cold water) and inversely proportional to the existing concentration of salt. This dissolution behaviour can be easily observed: salt dropped into a glass of cold water will dissolve slowly because the water surrounding the grains has a high concentration, and tends not to mix with the water above it. However, once stirred the salt dissolves immediately. The rate of dissolution is more sensitive to concentration than to temperature. For dry injection at medium to high flows, the dissolution occurs at the lower concentrations. As noted previously, dissolution is greatly enhanced by a good injection point such that even in glacier-fed streams where water temperatures are in the 1–3°C range, it can be assumed to occur instantaneously. For low flows,

and particularly in wide channels with limited turbulence, the salt can be dissolved in a bucket of water before injection to aid mixing. This does not alter the method as long as the salt mass is known and fully dissolved in the water.

After injection, the salt mixes into the stream by longitudinal dispersion, a process in which dissolved salt in the plume moves along its concentration gradient until a uniform concentration exists. The dispersion process is superimposed on the flow (hence the term “longitudinal”), which means that the plume extends downstream faster than it does towards the banks (Figure 4). In this figure, the line that



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Figure 4. Dispersion of salt at Russell Creek. Salt is injected at a point indicated by the arrow. The arcs represent the leading edge of the salt plume at successive time intervals as it disperses into the water column. The thickness of each line represents the salt concentration at each interval.

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is farthest downstream represents a state where the salt plume approaches full mixing, making point 1 in Figure 4 the best location for the conductivity probe. At any time during the measurement the salt concentration is higher at point 2 and lower at point 3 than it is at point 1. Therefore, locating the probe at point 2 will yield a quantity *A*, which is too high, resulting in an under-estimate of the flow. Similarly, if the probe is at point 3, *A* will be too low and the resulting flow estimate too high.

### Calibrating the Site: Mixing Length and Dosing Ratio

A conservative guideline is that the mixing length should be about 20 times the average channel width (Hudson and Fraser 2002). While Day (1977b) recommends mixing lengths of 25 times width (25XW), this is probably a conservative estimate since channel width is usually estimated visually. In most cases this can be assumed to be a “safe” mixing length. However, we have found that the optimum mixing length is often as low as 10 times the channel width, but users of the technique should conduct multiple trials to establish both the optimal injection points and mixing lengths for low and high flow at a particular site.

#### Determine the optimum mixing length

Since the rate of dispersion of the salt plume depends on several factors, each site will have a characteristic optimum mixing length. To determine this length, collect a sequence of measurements by varying the mixing length under stable flow conditions. The optimum mixing length is found where further increases in that length result in no relative change in the flow estimate.

#### Example of mixing length calibration

Culliton Creek is an ideal salt dilution site consisting of a long, straight reach with uniform gradient. The

morphology consists of a series of evenly spaced steps alternating with flow constrictions at each step. At

Discharge measurement	Probe and measurement	L (m)	Q (m <sup>3</sup> /s)
Q1	probe 2, M1	70	2.247
Q2	probe 1, M1	85	2.272
Q3	probe 2, M2	100	2.363
Q4	probe 1, M2	115	2.368

Culliton Creek two EC probes were used and two different injection points for 4 mixing lengths (Table 1; Figure 5). The channel is 10 m wide and the probes were deployed in mid-channel about 15 m apart. Two injection sites were chosen: the first site was about 70 m above the upper probe and the second about 30 m farther up the channel above the first injection site.

This resulted in 4 measurements at increasing mixing lengths (Table 1).

This procedure shows that the minimum mixing length at Culliton Creek is 100 m, or 10 times the channel width. A standard methodology therefore in determining the optimum *L* might involve varying the mixing length from 10X to 20X the average width in 10- to 15-m increments. For high water, the relatively high flow velocity that typifies steep channels suggests that the mixing length should default to 20X the channel width.

#### Determine optimum salt dosing

To apply dry injection for flow measurement, the injection mass

must be known. This requires an accurate scale and a means of packaging salt for use in the field. Zip-lock freezer bags can easily hold up to 2 kg of salt. Bulk salt (usually obtained in 20-kg bags from a bulk food distributor) can be pre-weighed into packages of various masses that can be made up in the field to the desired amount. These bags can be carried to remote sites in a backpack. Empty bags are weighed upon return to the lab or office to account for any residual salt not injected.

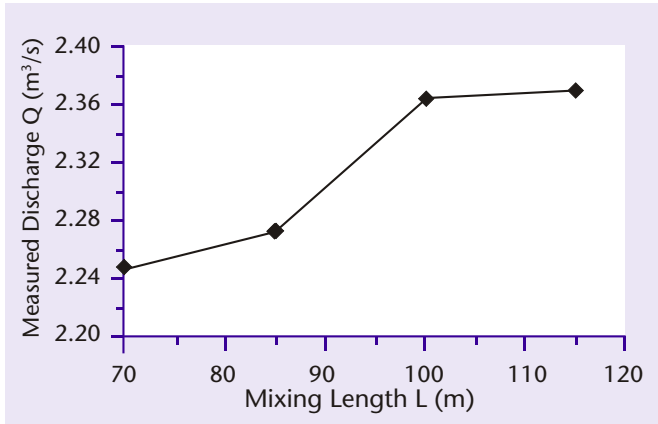


Figure 5. This chart shows how the increase in mixing length between each flow measurement results in an increase in measured flow to a maximum. The 100-m L is the optimum L.

The final step to determine salt dosing for a new site is to adjust the dosing mass needed to get a clean signal. In an earlier report (Hudson and Fraser 2002) we recommended dosing at the rate of 2 kg/m<sup>3</sup>/s of flow (dosing ratio = 2). Since then, we have found that in many cases the dosing ratio can be as low as 0.5. The aim of salt dosing is to add enough salt to get a clean signal without exceeding the most sensitive toxicity threshold (Moore 2004a, 2004b) of 400 mg/L NOEC (no observed effect concentration) for *Rana breviceps* (frog).

To get a strong signal, the difference between baseline and peak EC should be at least 100 times the resolution of the probe. Optimum dosing ratio is related to the optimal mixing length;



**Table 2. A first approximation of dosing ratio in relation to the optimal mixing length (L)**

Mixing length L (m)	Dosing ratio
10XW	0.5 – 1.0
15XW	1.5 – 2.5
20XW	2.0 – 3.0

the longer the mixing length, the more salt is needed. At a mixing length of 10XW, a dosing ratio of 1.0 works well. For longer mixing lengths, more salt is often needed for a clean EC signal (Table 2).

### Limitations

Understanding the limits of applicability of dilution gauging methods in general will help to determine the appropriate technique for a given gauging site and whether to trust individual measurements. Operators should document their observations, thus contributing to a systematic assessment of the limitations of any technique and leading to informed decisions about their choice.

Here are some of our observations regarding the limits of applicability of the mass balance method:

1. Instream conditions: Turbulence is all-important. As a rule of thumb, conditions that violate the assumptions of current metering favour salt dilution and vice-versa. When applying the mass balance method with dry salt, try to observe the salt as it is injected. If it falls to the bottom of the channel and sits there in clumps, dry injection will not work. Either dissolve the salt or use constant rate injection or current metering. Some channel conditions render any method of measurement difficult or impossible. These include low flow conditions in very wide channels where the flow is dispersed by channel sediment.
2. There are situations where injected salt may be absorbed by (or may adsorb to) vegetation in

the channel or other substances (e.g., neoprene chest waders). Try to avoid applying salt dilution in channel reaches with a lot of aquatic vegetation. Always inject the salt in a downstream direction and, if possible, keep out of the channel during a measurement.

3. Violation of assumptions. In practice it is difficult not to violate some assumptions such as pools in the gauging reach. In riffle-pool channels, this is problematic since violation is a matter of degree — it can be minimized but seldom eliminated. If pools are large relative to channel area, then the salt will hang in the pool resulting in an extended tail. In some cases, pooling results in less than 5% error due to cutting off the tail of the distribution (Hudson and Fraser 2002) but the error will be systematic (i.e., it will tend to give an over-estimate).

4. Examine the conductivity-over-time graph. A clean dump with adequate dosing and full mixing has a characteristic shape. Many common problems can be detected this way:

- a. A smooth graph with a strong peak and short tail indicates a good chance of success.

b. Irregularities in the graph could indicate low dosing, improper mixing, or other problems that could render a measurement unreliable. For example:

- i. Double peak — is it lack of mixing or is it discontinuous injection?
- ii. Extended tail — does it indicate pooling of salt or a changing baseline?

This list of limitations is not comprehensive. The operator should recognize that documentation of observed limitations will lead to improved confidence in the application of dilution methods. There is no substitute for experience in applying salt dilution gauging. Over time an experienced operator will be able to judge the applicability of the method to a given site.

### Comparison of Solution Injection and Mass Balance (Dry Injection) Methods

The mass balance method and slug injection of salt in solution method require similar calculations and both possess similar requirements for selecting a suitable measurement reach. Both methods also require full lateral mixing of the salt in the stream

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**Table 3. Comparison of requirements of solution and dry injection methods**

	Salt solution method	Dry injection (mass balance) method
Equipment requirements (not including EC logger)	Volumetric measurement equipment - 13 items (Moore 2005)	Weigh scale & packaging A bucket or similar for injecting Calibration kit (4 items)
Upper limits	Limited ability to measure high flow (upper limit 20–25 m <sup>3</sup> /s)	More-or-less full range of flows for small watersheds (i.e., up to 100 m <sup>3</sup> /s)
Calibration	Calibration of injection solution required for each measurement	Simple procedure to calibrate probe or to verify existing calibration. Required at each measurement for QA/QC
Low flow	Applicability depends on the channel morphology — for wide or braided channels at low flow, default to constant injection method (Moore 2004b)	Same as solution injection by dissolving the salt mass in streamwater before injection



and measurement of conductivity over time at a point downstream.

The mass balance method was developed for ease of application in the field. Solution injection with saturated (20%) solution requires the operator to manipulate a slug that is approximately five times more massive than a dry salt slug. In practice, a saturated solution is difficult to create in the field — a solution of 10–15% is more realistic. Using 0.5 kg of salt per cubic metre per second as a guideline to assess the upper limit of applicability and if the flow is 10 m<sup>3</sup>/s, dry salt will require the operator to inject at least 5 kg of salt while solution injection will require about 35 L, or approximately 35 kg, of solution to be injected. While this amount may be manageable, the ratio makes dry salt injection more appealing at higher flows. For example, we have measured flows greater than 40 m<sup>3</sup>/s with dry injection. To use solution to measure the same flow would require at least 300 L of solution. Thus the practical upper limit of applicability of the salt dilution method by solution is in the range of 5–10 m<sup>3</sup>/s. The use of dry salt instead of salt solution allows the upper limit to be extended to 100 m<sup>3</sup>/s or more.

Another advantage of the mass balance method is that it requires less equipment than the solution injection method (Table 3). The solution method requires the creation and calibration of injection solution for each measurement. This requires more time on site and more equipment to be carried to the field as well as more opportunity for error in calibration.

Salt solution has an advantage over dry injection in that it will mix more readily in lower, less turbulent flow, since the salt is already dissolved in water. However, the mass balance method can also be used with solution as a known mass of salt can be dissolved in a bucket of

streamwater before injection and the calculation carried out as though it were a dry mass.

## Summary/Conclusions

Whether using solution or dry injection methods, the operator needs to have equipment for measurement and calibration. As long as the application criteria are met, it makes little difference which method is used. Both methods are capable of high precision. We have tested the mass balance method with dry injection under a wide range of conditions. Its speed and simplicity of application in the field make it an operational standard hydrometric method for small, steep streams.

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### For more information, contact:

**Rob Hudson, P.Geo., Ph.D., Hydrologist**

613 Redwood Drive  
Qualicum Beach, BC  
V9K 1A2

**Tel: (604) 341-0135**

**E-mail: Robert.Hudson@telus.net**

**John Fraser**

Fraser Hydro-Technical  
**E-mail: Fras@telus.net**

### References

- Church, M. 1975. *Electrochemical and fluorometric tracer techniques for streamflow measurements*. British Geomorphological Research Group. Technical Bulletin #12.
- Church, M. and R. Kellerhals. 1970. *Stream gauging techniques for remote areas*

*using portable equipment*. Department of Energy, Mines and Resources Canada, Ottawa, Ont., Technical Bulletin #25, pp. 55–68.

- Collins, D.N. 1978. *Hydrology of an alpine glacier as indicated by the geochemical composition of meltwater*. *Zeitschrift für Gletscherkunde und Glazialgeologie* 13(1/2):219–238.
- Day, T.J. 1976. *On the precision of salt dilution gauging*. *Journal of Hydrology* 31:293–306.
- Day, T.J. 1977a. *Field procedures and evaluation of a slug dilution gauging method in mountain streams*. *Journal of Hydrology (NZ)* 16(2):113–133.
- Day, T.J. 1977b. *Observed mixing lengths in mountain streams*. *Journal of Hydrology* 35:125–136.
- Elder, K., R. Kattelmann, and R. Ferguson. 1990. *Refinements in dilution gauging for mountain streams*. In *Hydrology in Mountainous Regions. I – Hydrological Measurements; the Water Cycle*. International Association for Hydrological Science Proceedings of two Lausanne symposia, August 1990. IAHS Publication No. 193, pp. 247–254.
- Hongve, D. 1987. *A revised procedure for discharge measurement by means of the salt dilution method*. *Hydrological Processes* 1:267–270.
- Hudson, R. and J. Fraser. 2002. *Alternative methods of flow rating in small coastal streams*. B.C. Ministry of Forests, Vancouver Forest Region, Nanaimo, BC. Extension Note EN-014 Hydrology. 11 p.
- Johnstone, D.E. 1988. *Some recent developments of constant-rate salt dilution gauging in rivers*. *Journal of Hydrology (NZ)* 27:128–153.
- Kite, G. 1993. *Computerized streamflow measurement using slug injection*. *Hydrological Processes* 7:227–233.
- Moore, R.D. 2004a. *Introduction to salt dilution gauging for streamflow measurement: Part 1*. *Streamline Watershed Management Bulletin* 7(4):20–23.
- Moore, R.D. 2004b. *Introduction to salt dilution gauging for streamflow measurement Part II: Constant-rate injection*. *Streamline Watershed Management Bulletin* 8(1):11–15.
- Moore, R.D. 2005. *Introduction to salt dilution gauging for streamflow measurement Part III: Slug injection using salt in solution*. *Streamline Watershed Management Bulletin* 8(2):1–6.
- Østrem, G. 1964. *A method of measuring water discharge in turbulent streams*. *Geographical Bulletin* 21:21–43.
- Smart, C.C. 1992. *Temperature compensation of electrical conductivity in glacial meltwaters*. *Journal of Glaciology* 38:9–12.