


**INVESTIGATING THE TIMING OF DEGLACIATION AND  
THE EFFICIENCY OF SUBGLACIAL EROSION IN  
CENTRAL-WESTERN GREENLAND WITH COSMOGENIC  
 $^{10}\text{Be}$  AND  $^{26}\text{Al}$**





Joseph Graly  
University of Vermont '10  
University of Wyoming

Tom Neumann  
NASA Cryospheric Branch

Paul Bierman  
University of Vermont

Bob Finkel  
Lawrence Livermore National Laboratory



## Thesis Road Map:

OPTIMIZING SAMPLE PREPARATION FOR  
 $^{10}\text{Be}$  ANALYSIS BY ACCELERATOR MASS  
SPECTROMETRY

PAIRED BEDROCK AND BOULDER  $^{10}\text{Be}$   
CONCENTRATIONS RESULTING FROM  
EARLY HOLOCENE ICE RETREAT NEAR  
JAKOBHAVN ISFJORD, WESTERN  
GREENLAND

CONSTRAINING LANDSCAPE HISTORY  
WITH  $^{10}\text{Be}$  AND  $^{26}\text{Al}$  IN PAIRED BEDROCK  
AND BOULDER SAMPLES, UPERNAVIK,  
CENTRAL-WESTERN GREENLAND



Background  
&  
Rationale

# Why Greenland?

- Greenland Ice Sheet holds ~7 m global sea level equivalent
- Highly susceptible to warming climate

Goal: to investigate how the ice sheet behaved during *past* warming episodes in order to understand how it might behave in the future



# Climate Basics

Climate is not static over time!

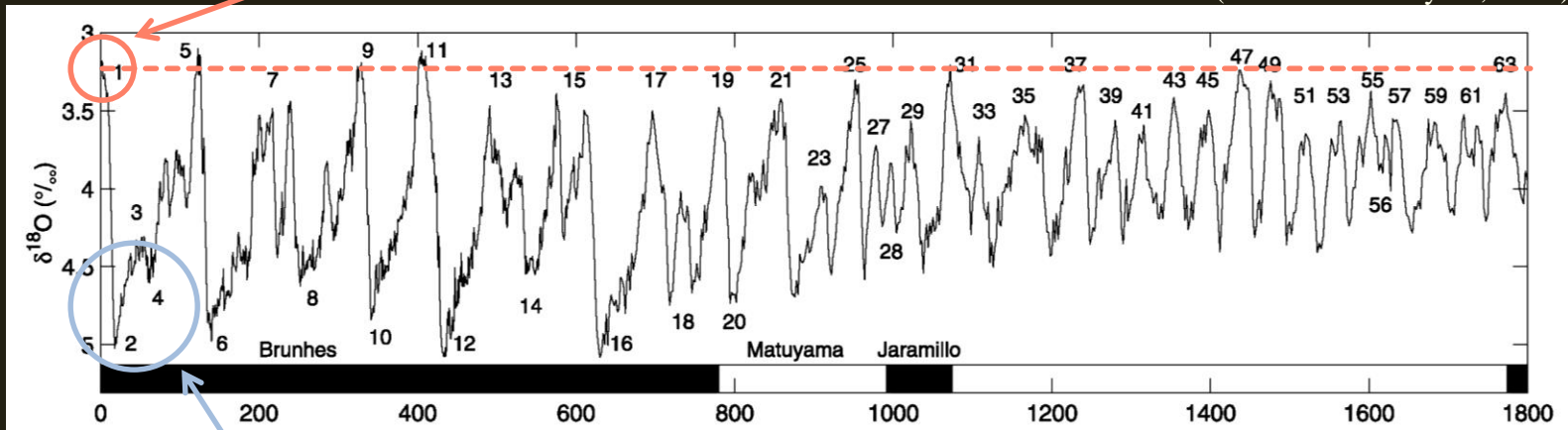
Holocene Period

(Lisiecki and Raymo, 2005)

Warmer



Colder



X-axis: thousands of years ago (“ka”)

Last “Ice Age”

Warm periods: “Interglacial”

Cold periods: “Glacial”

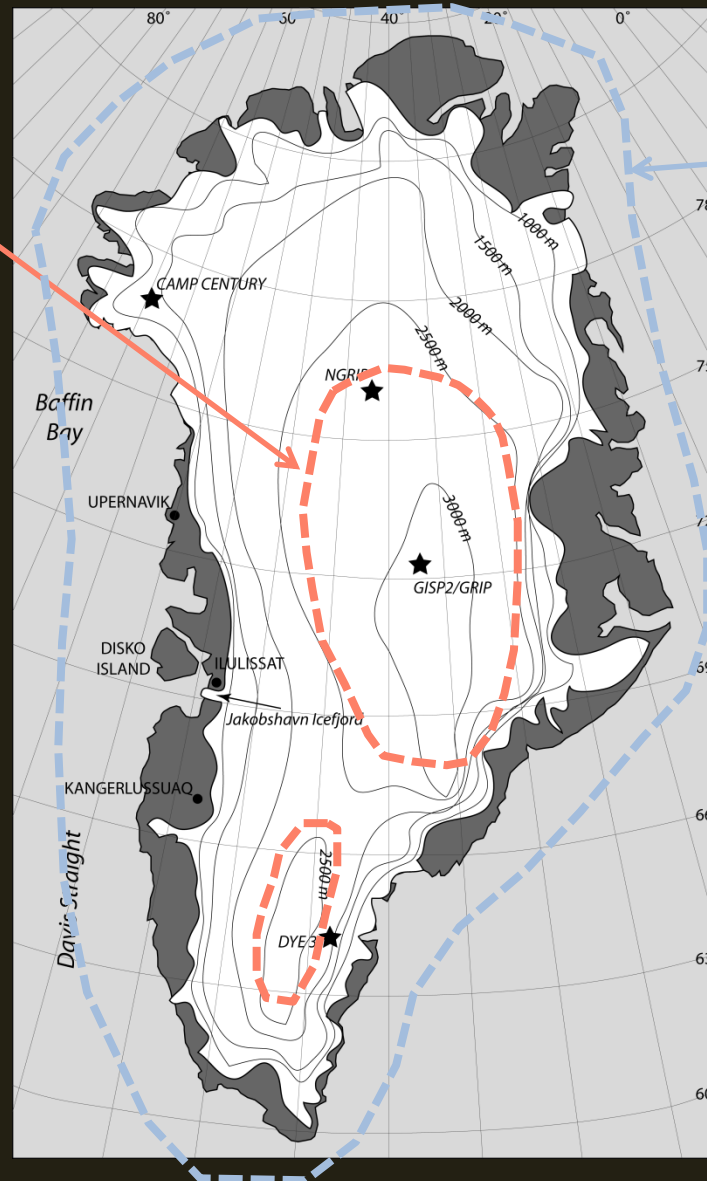
# Variability of the Greenland Ice Sheet

Interglacial period  
ice extent

Glacial period  
ice extent

Cuffey and Marshall (2000)  
Létrégilly et al. (1991)  
Otto-Bliesner et al. (2006)  
Overpeck et al. (2006)

Bennike et al. (2002)  
Funder and Hansen (1996)  
Winkelman et al (2010)



# Ice Can Lose Mass in Multiple Ways



Melting

Sublimation



Calving





# Research Goals

At two different sites in western Greenland:

- Make inferences about the efficiency of subglacial erosion
  - How effectively does ice erode bedrock surfaces?
  - Does this control the landscape we see today?
- Determine the chronology of ice retreat after the last glacial period
  - When did ice retreat begin?
  - How long did ice retreat last?
  - How rapid were ice retreat rates?
- Compare ice behavior between two sites





Tools,  
Study Design,  
&  
Methodology

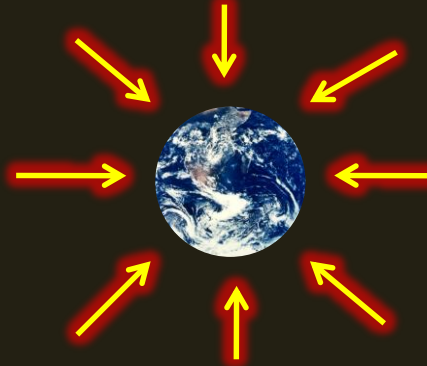
# Tools: *in situ* cosmogenic $^{10}\text{Be}$ and $^{26}\text{Al}$

- “*In situ*”: produced within the mineral structure (quartz)
- “Cosmogenic”: from cosmic rays
- “ $^{10}\text{Be}$ ”: rare, radioactive isotope of beryllium
- “ $^{26}\text{Al}$ ”: rare, radioactive isotope of aluminum

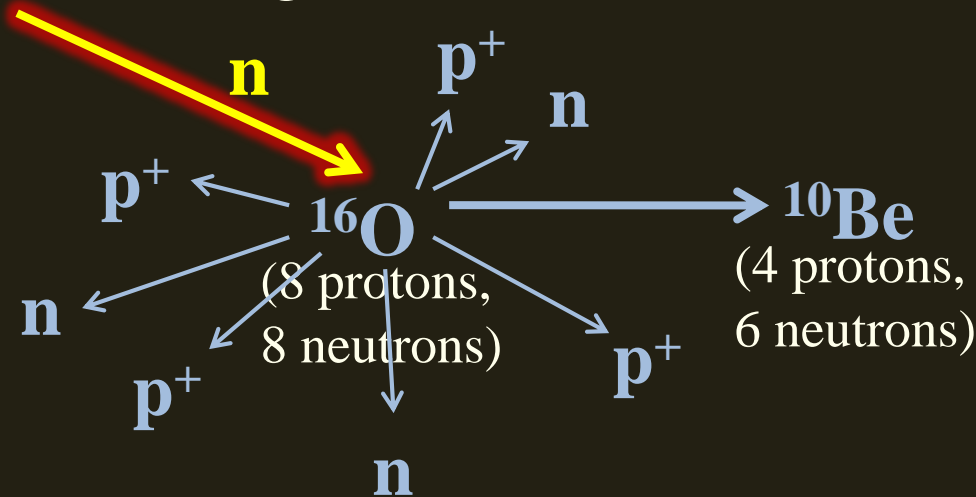


# Formation of *in situ* cosmogenic $^{10}\text{Be}$

Earth is bombarded by high-energy cosmic rays



...causing the formation of  $^{10}\text{Be}$  in quartz



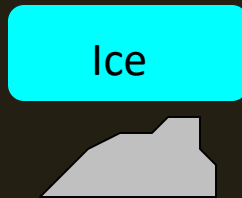
$^{10}\text{Be}$  is produced only on the surface of a rock

$^{10}\text{Be}$  is produced at about 6 atoms per year per gram of quartz

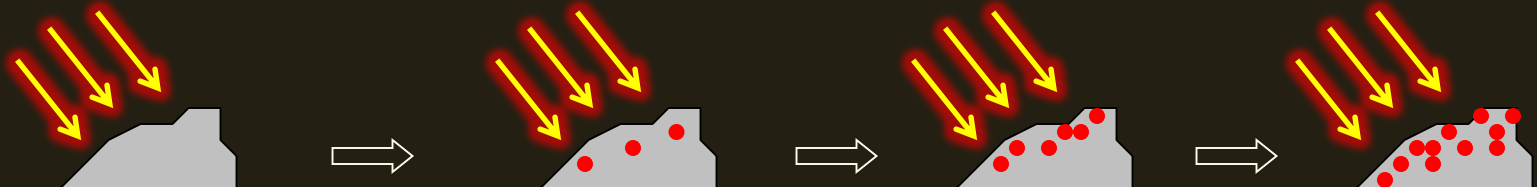
$^{10}\text{Be}$  is radioactive and has a half-life of 1.38 million years

# “Cosmogenic Dating”

Glacial period: Bedrock is **shielded**



Interglacial period: Bedrock is **exposed**



Assumption: Zero **inheritance** (i.e. no  $^{10}\text{Be}$  leftover from previous periods of exposure)

# Sampling Scheme & Methodology

- Collect bedrock and boulder samples in a transect parallel to direction of ice flow
- Analyze  $^{10}\text{Be}$  concentrations
  - Isolate quartz, remove impurities
  - Isolate pure Be from quartz
  - Measure  $^{10}\text{Be}/^9\text{Be}$  ratios by accelerator mass spectrometry (AMS)
  - Calculate exposure ages
- Analyze  $^{26}\text{Al}$  contents (only certain samples)
- Make inferences about ice behavior

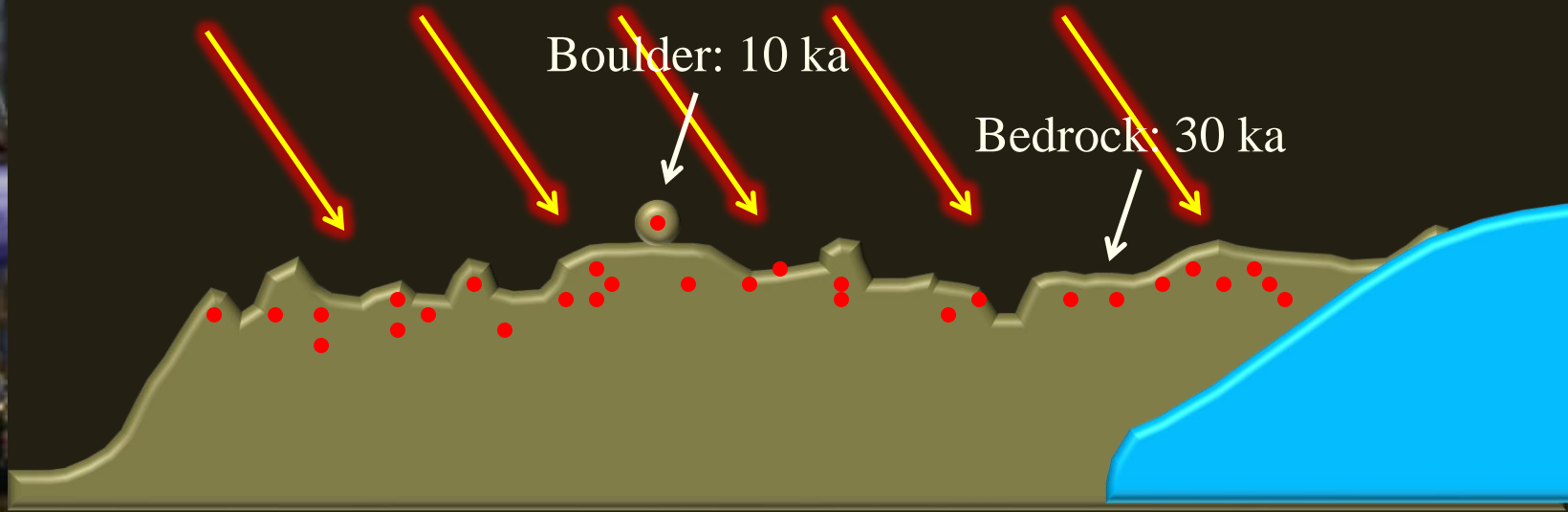
# Sampling Scheme

“Dipstick” Sampling:



# Cosmogenic Inheritance

- More prevalent in bedrock than in boulders
  - Outcrop is exposed (earlier interglacial?)
  - Outcrop is covered by ice
  - Ice is non-erosive, doesn't remove  $^{10}\text{Be}$
  - Outcrop is exposed again





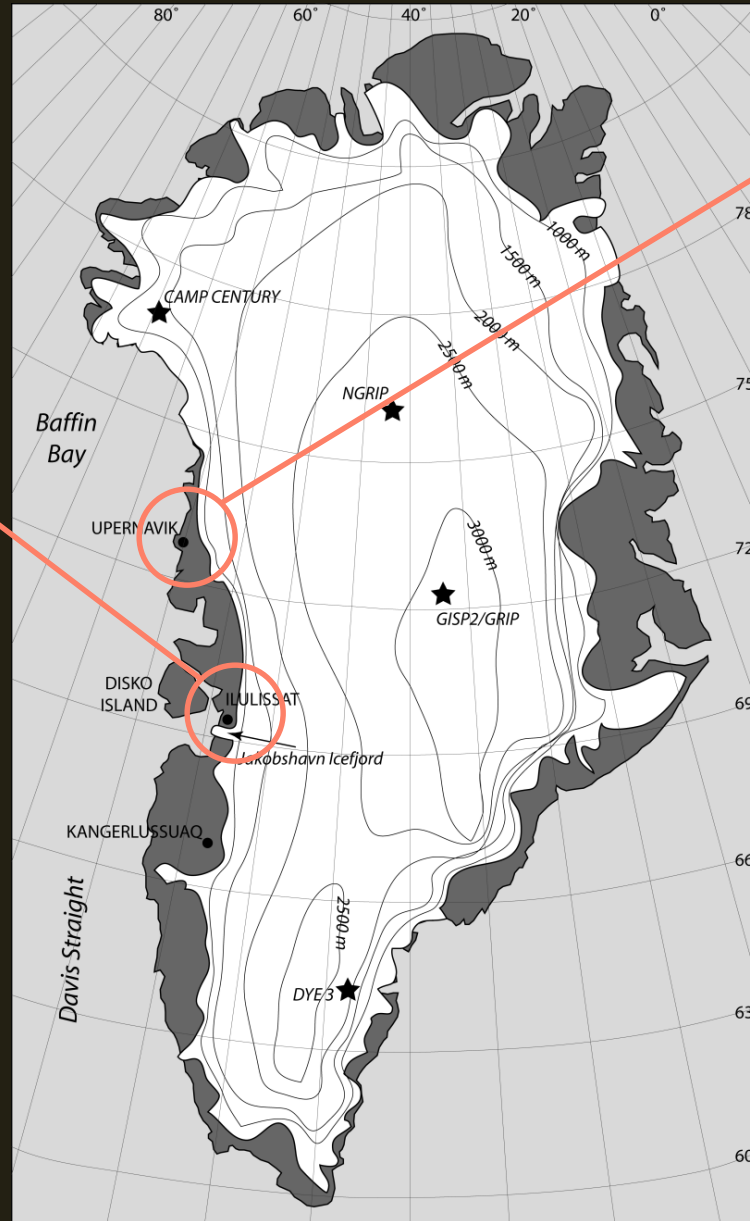
# Study Sites

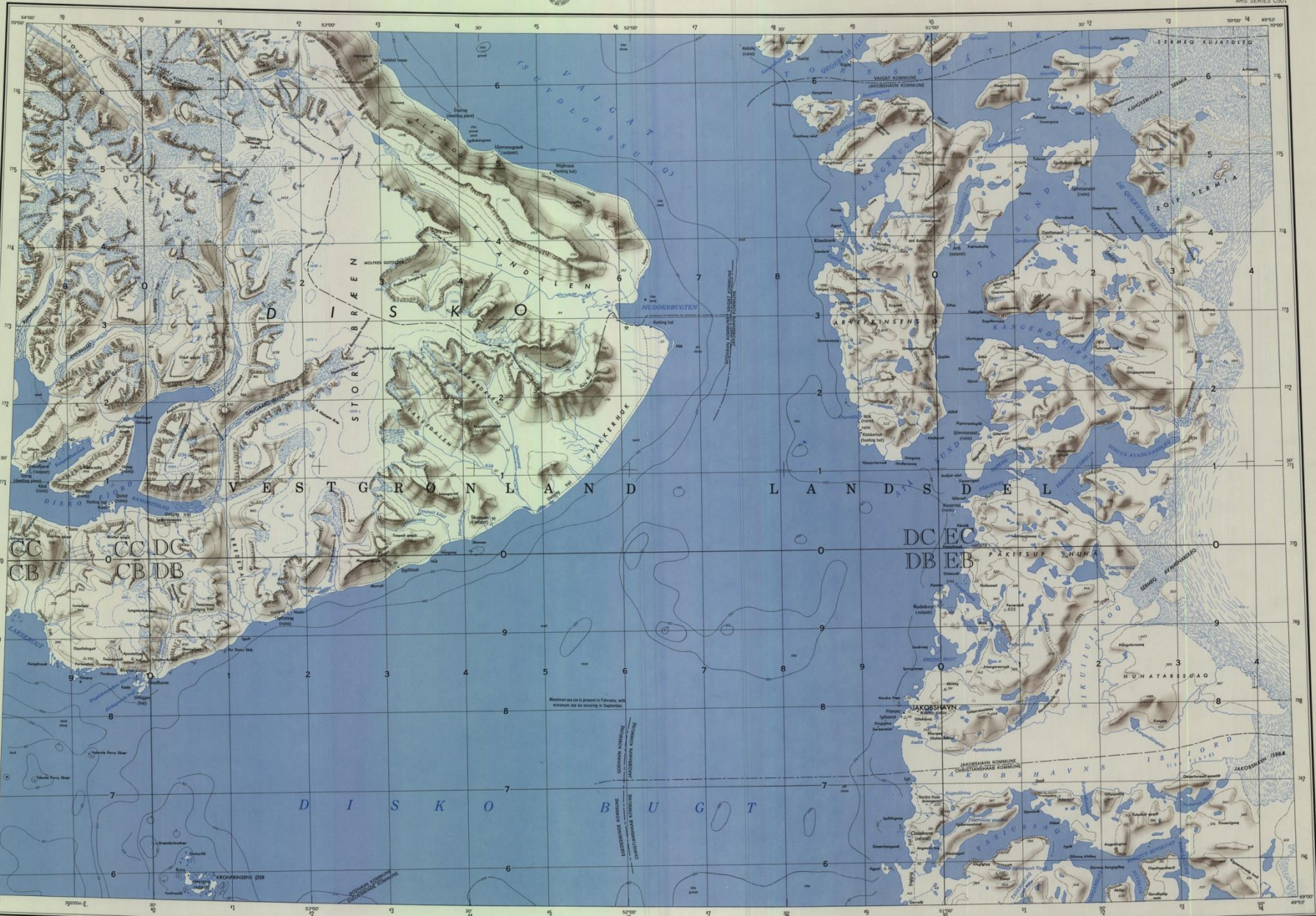
## Iulissat:

- Latitude: 69°N
- Continuous land surface
- Jakobshavn Isfjord

## Upernavik:

- Latitude: 73°N
- Fjord-dissected terrain
- No major outlet glaciers





AMS C501  
Edition 1-AMS (First Printing, 4-65)  
Prepared by the Army Map Service (AMS), Corps of Engineers, 13 S. Ave., Washington, D. C. Compiled in 1964 from: Geovisat, 1:250,000, Geografiska Institutet, Sweden (D. F. 1, 10 N. 1, 1:50,000); and Geovisat, 1:250,000, Geografiska Institutet, Sweden (D. F. 1, 10 N. 1, 1:50,000). (1:50,000) and Geovisat, 1:250,000, Geografiska Institutet, Sweden (D. F. 1, 10 N. 1, 1:50,000).  
Name data are taken directly from Danish sources, except that a Q has been used for a K or a K.

POPULATED PLACES

Over 1,000  
750 to 1,000  
500 to 750  
250 to 500  
Less than 250

**GLOSSARY**

Ship, beam	gullet	harbor	glacier
Flag, height	high, crest	farm, wall	glacier
Line, depth	valley	island	glacier
Low, depth	water	island	glacier
Low, depth	water	island	glacier
Low, depth	water	island	glacier
Low, depth	water	island	glacier
Low, depth	water	island	glacier
Low, depth	water	island	glacier

**LEGEND**

**ROADS**

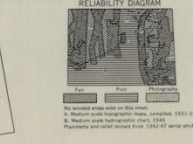
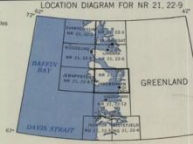
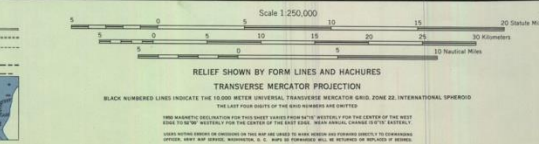
- Loose surface, graded, all weather
- Loose surface, no weather, no dirt
- Trail

**BOUNDARIES**

- Land-Sea
- Continental

**Landmark Features**

- Water: Beach, Bank
- Wreck: Exposed, Sunken
- Manicled control point
- Altimeter position
- Seaplane anchorage
- Shoal
- Exclosure Pier
- Depth curves in fathoms
- Star
- Star: Resurvey



PRINTED BY ARMY MAP SERVICE, CORPS OF ENGINEERS, 4-96, 374001

THIS IS A TECHNICAL SHEET OF THE 1:250,000 SCALE SERIES C501.

**CC DC EC**  
**CB DB EB**

1. Red letters identifying 100,000 scale sheets to which this sheet is adjacent. To locate this sheet, go to the left of and above the letter designating the sheet on the grid. The sheet number on the grid will be the sheet number of the sheet to which this sheet is adjacent.

2. Green letters identifying 1:250,000 scale sheets to which this sheet is adjacent. To locate this sheet, go to the left of and above the letter designating the sheet on the grid. The sheet number on the grid will be the sheet number of the sheet to which this sheet is adjacent.

3. Black letters identifying 1:50,000 scale sheets to which this sheet is adjacent. To locate this sheet, go to the left of and above the letter designating the sheet on the grid. The sheet number on the grid will be the sheet number of the sheet to which this sheet is adjacent.

4. Red letters identifying 1:100,000 scale sheets to which this sheet is adjacent. To locate this sheet, go to the left of and above the letter designating the sheet on the grid. The sheet number on the grid will be the sheet number of the sheet to which this sheet is adjacent.

5. Black letters identifying 1:50,000 scale sheets to which this sheet is adjacent. To locate this sheet, go to the left of and above the letter designating the sheet on the grid. The sheet number on the grid will be the sheet number of the sheet to which this sheet is adjacent.

6. Red letters identifying 1:100,000 scale sheets to which this sheet is adjacent. To locate this sheet, go to the left of and above the letter designating the sheet on the grid. The sheet number on the grid will be the sheet number of the sheet to which this sheet is adjacent.

7. Black letters identifying 1:50,000 scale sheets to which this sheet is adjacent. To locate this sheet, go to the left of and above the letter designating the sheet on the grid. The sheet number on the grid will be the sheet number of the sheet to which this sheet is adjacent.

8. Red letters identifying 1:100,000 scale sheets to which this sheet is adjacent. To locate this sheet, go to the left of and above the letter designating the sheet on the grid. The sheet number on the grid will be the sheet number of the sheet to which this sheet is adjacent.

9. Black letters identifying 1:50,000 scale sheets to which this sheet is adjacent. To locate this sheet, go to the left of and above the letter designating the sheet on the grid. The sheet number on the grid will be the sheet number of the sheet to which this sheet is adjacent.

10. Red letters identifying 1:100,000 scale sheets to which this sheet is adjacent. To locate this sheet, go to the left of and above the letter designating the sheet on the grid. The sheet number on the grid will be the sheet number of the sheet to which this sheet is adjacent.

Site #1: *Ilulissat, 69°N*



Site #1: *Ilulissat, 69°N*



Site #1: *Ilulissat, 69°N*





# Site #2: *Upernavik, 73 °N*



Site #2: *Upernavik, 73 °N*





# Site #2: *Upernavik, 73 °N*



# Recap

- $^{10}\text{Be}$  and  $^{26}\text{Al}$  are radioactive isotopes formed when cosmic rays interact with quartz
- Use the production rate and concentration to calculate an exposure age
- Two study sites in western Greenland: Ilulissat ( $69^\circ\text{N}$ ) and Upernavik ( $73^\circ\text{N}$ )
- Samples collected in “dipsticks”
- Goal: to understand subglacial erosion efficiency and ice retreat characteristics



# Thesis Part I: Methodological Development

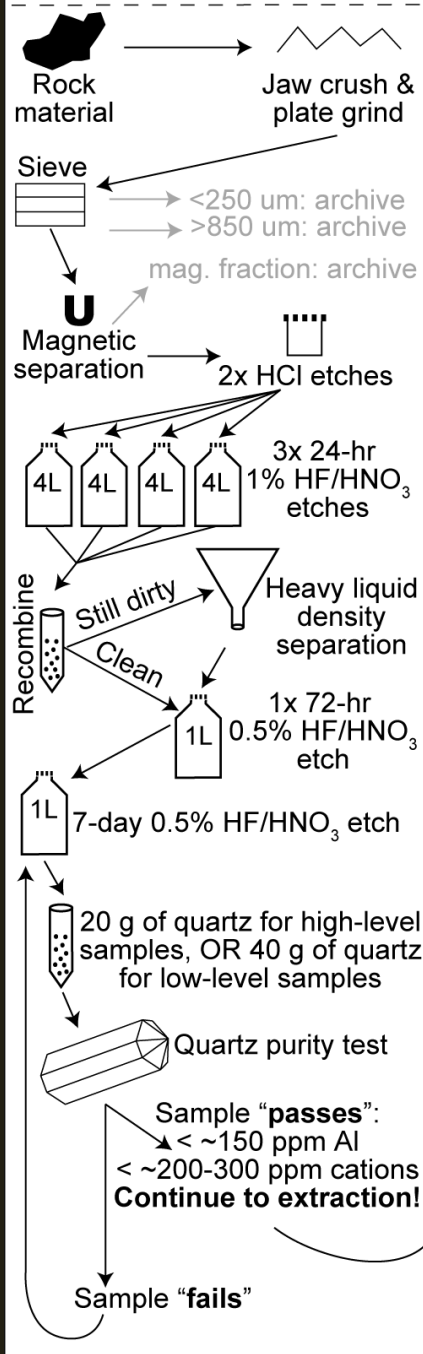
# Methodological Limitations

- Counting individual atoms!
- Limited by the number of atoms that can be counted in a given amount of time
- More atoms counted = higher precision
- AMS counting efficiency is controlled by the “**beam current**”
  - Purity of sample
  - Amount of sample
- Higher beam currents = higher counting efficiency = higher precision

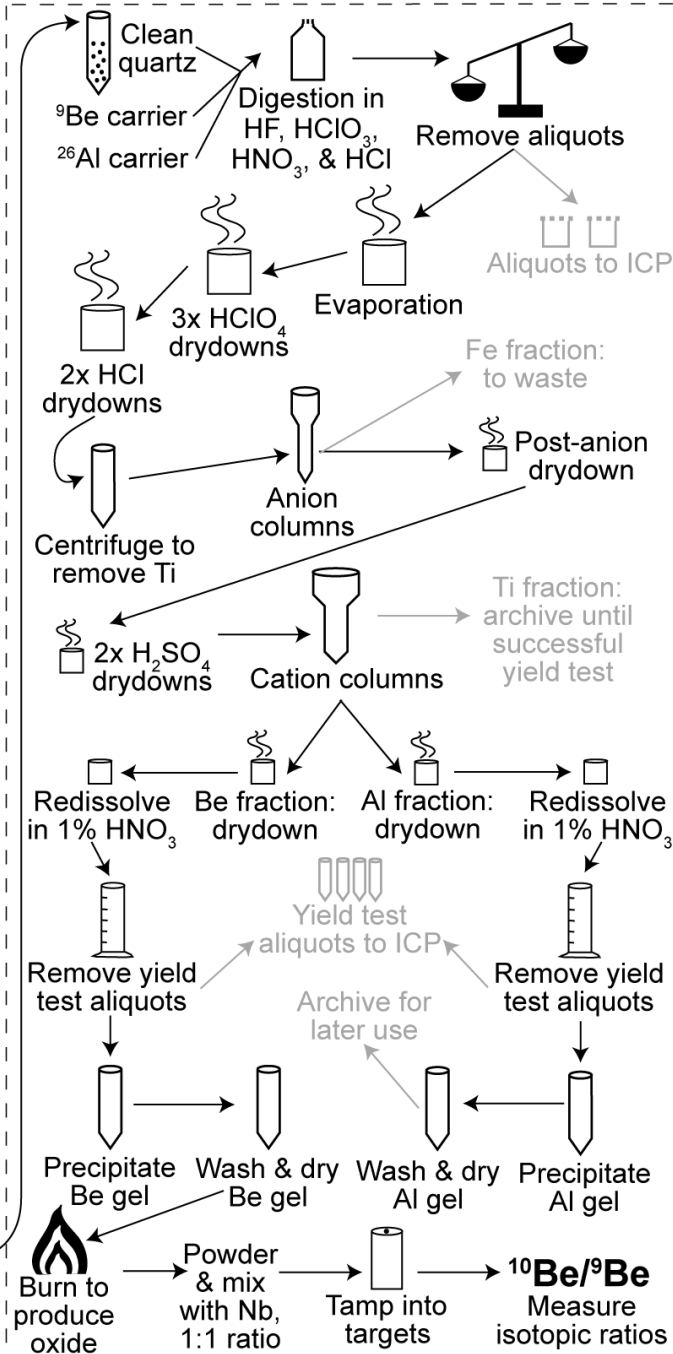


# Laboratory Methods

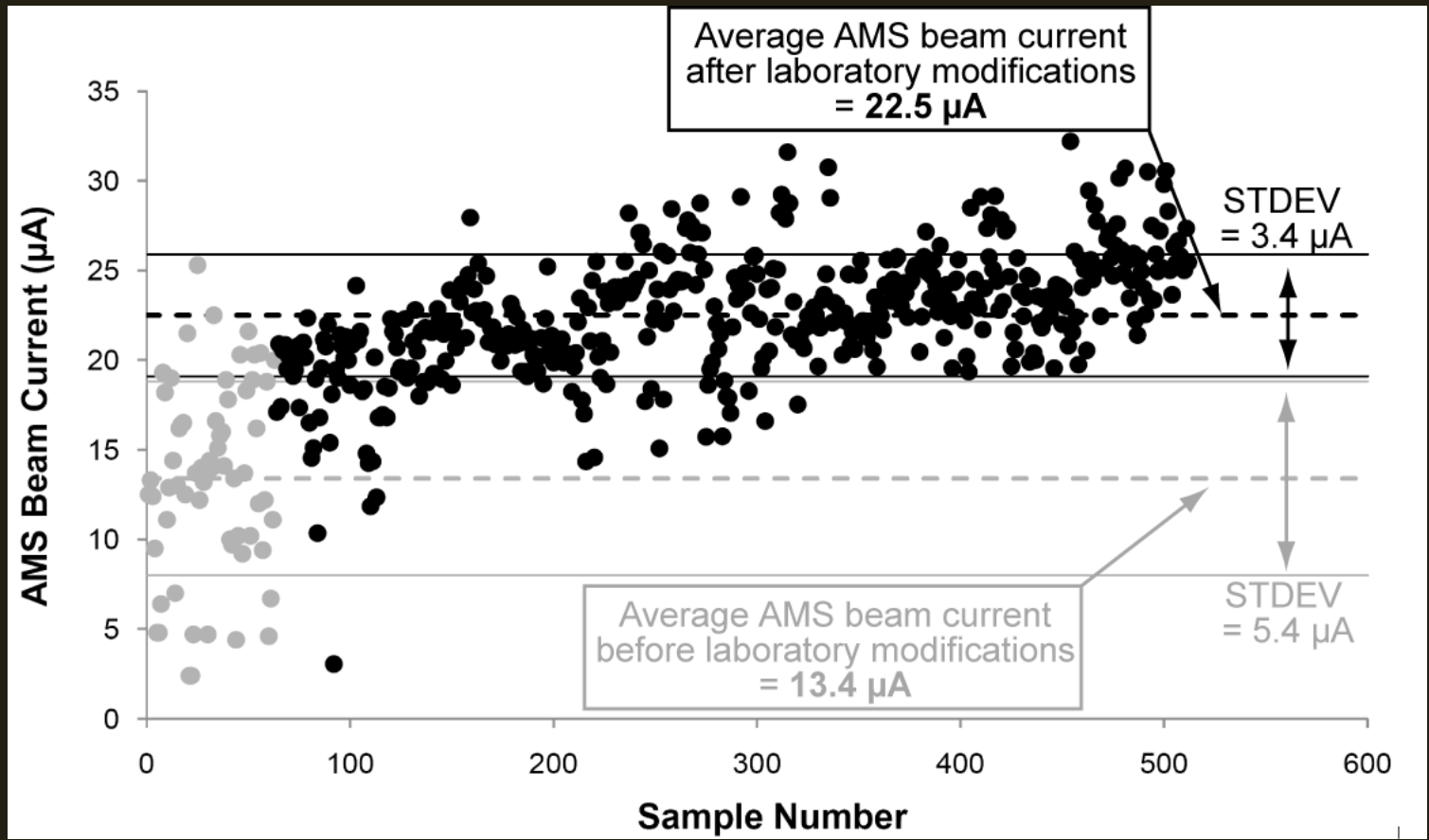
## Mineral Separation Lab Quartz Purification



## Cosmogenic Lab Be and Al Extraction

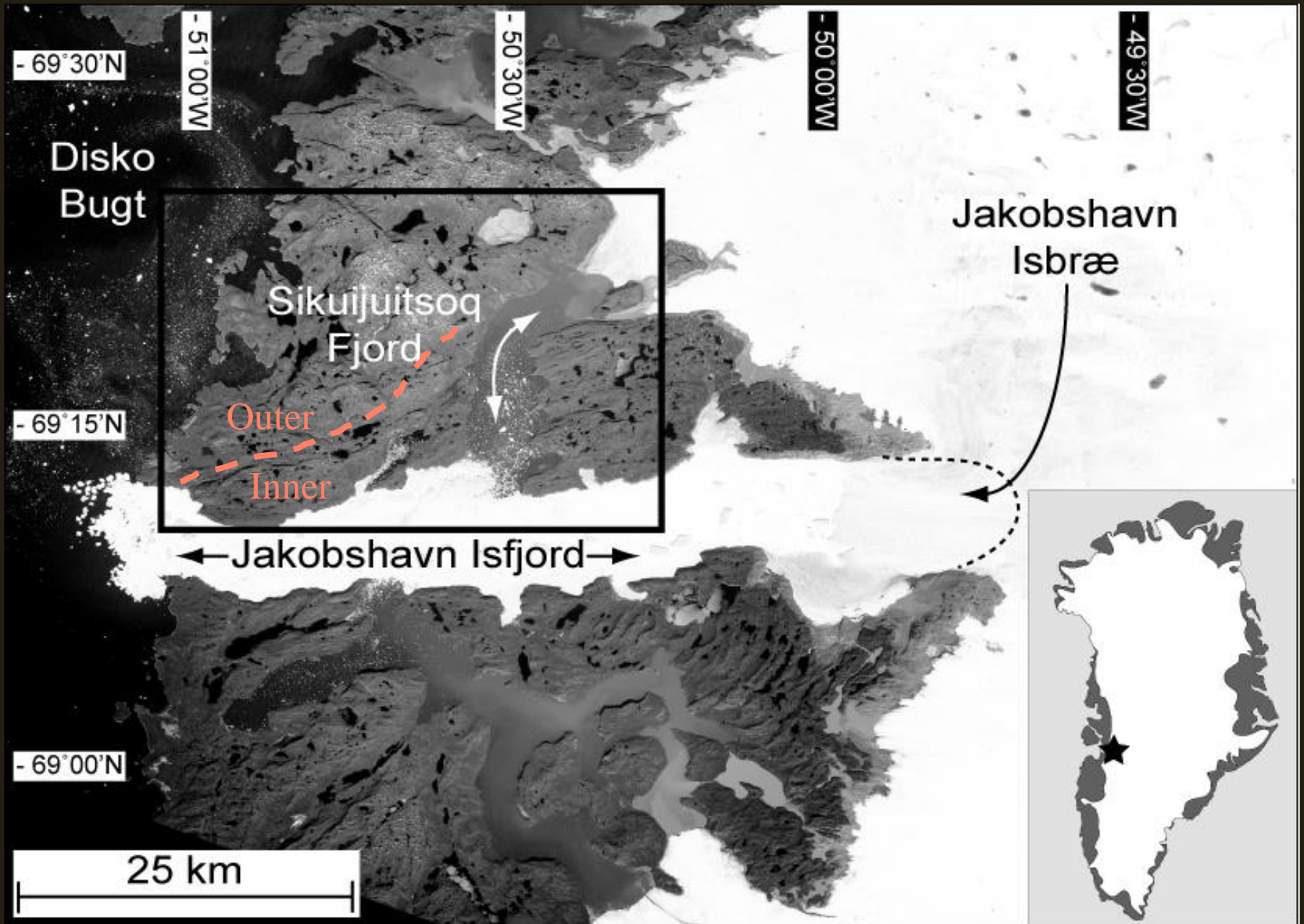


# Methodological Optimization





# Thesis Part II: Ilulissat, Greenland



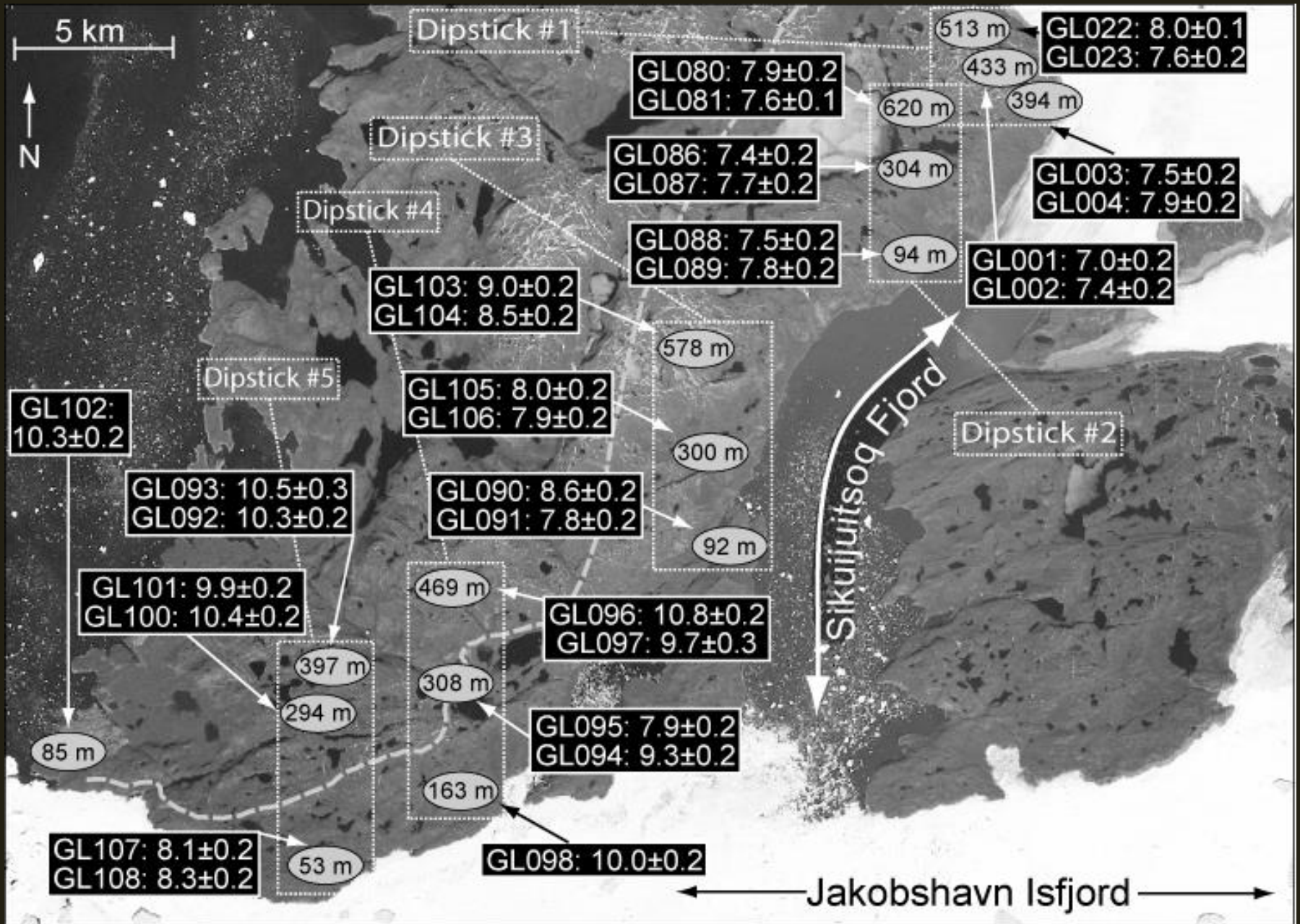
(Image courtesy of Landsat, 2000)



# Fjord Stade Moraine



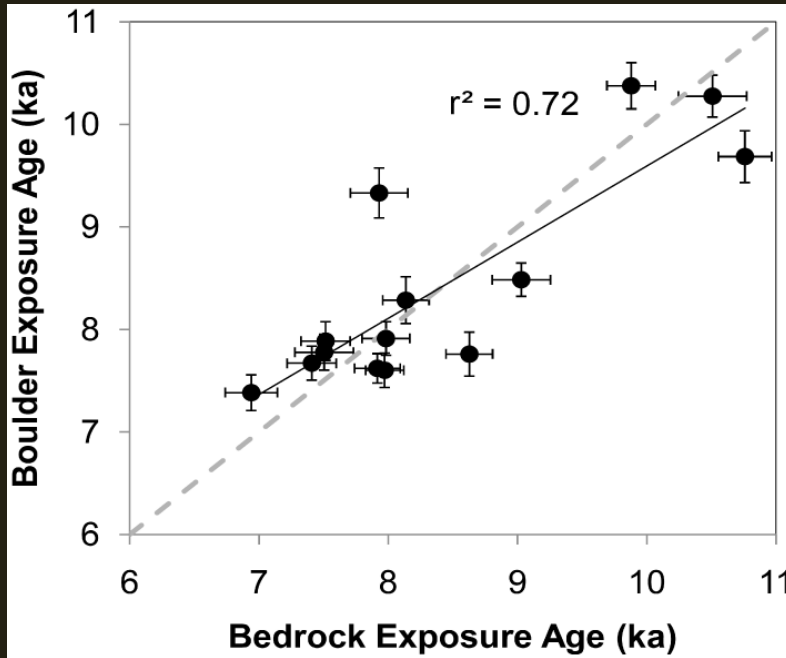
(Photo courtesy of Nicolas Young)



(Image courtesy of Landsat, 2000)

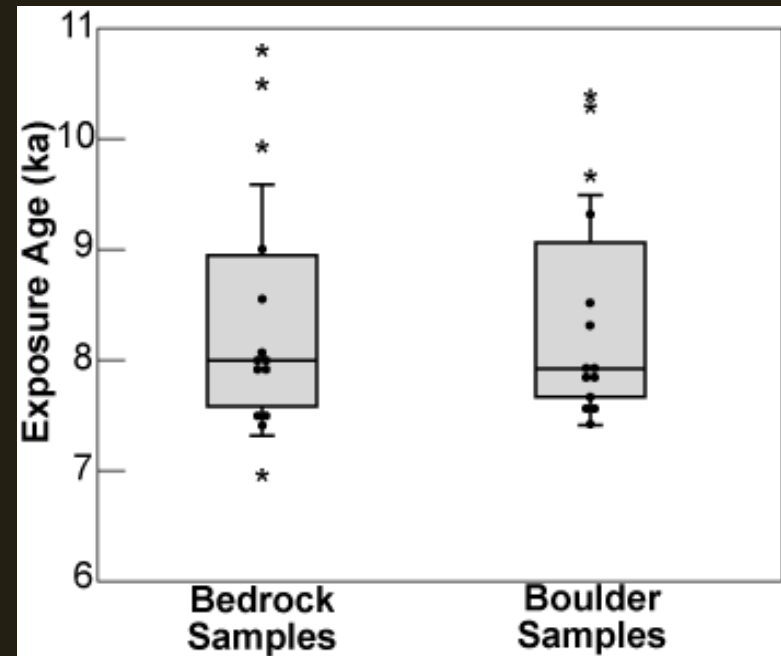
All ages are Holocene

# Bedrock/Boulder Comparison

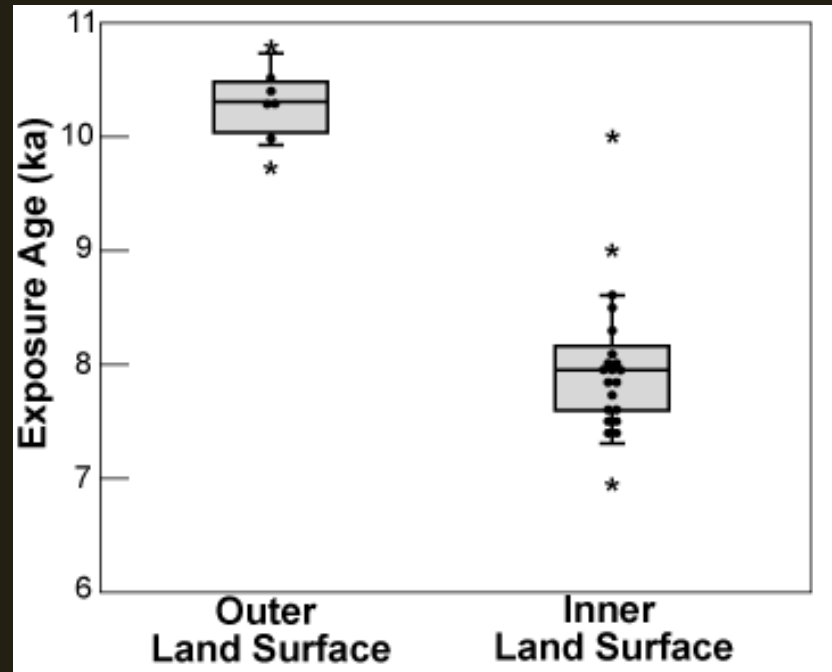


Bedrock and boulder samples are in close agreement

Little or no inheritance  
Erosive glacial ice



# Complicated Deglaciation Pattern



Age of outer land surface:  $10.3 \pm 0.4$  ka (n = 7)

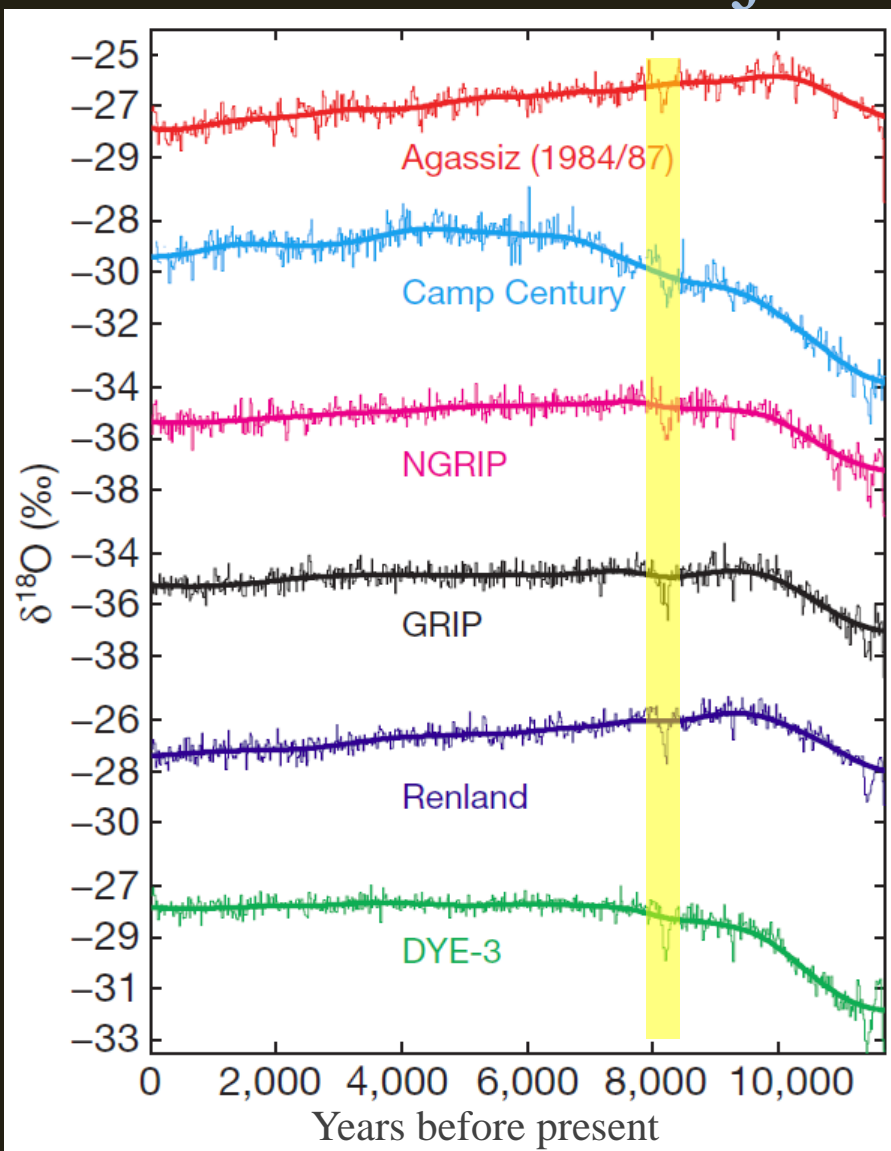
Age of inner land surface:  $8.0 \pm 0.7$  ka (n = 21)

Just inside moraine:  $8.2 \pm 0.1$  ka (n = 2)

**Complicated deglaciation pattern!**

**Age of the Fjord Stade moraine is  $\sim 8.2$  ka**

# Formation of the Fjord Stade Moraine



Fjord Stade moraine formed due to an ice margin re-advance in association with the “8200 Event”

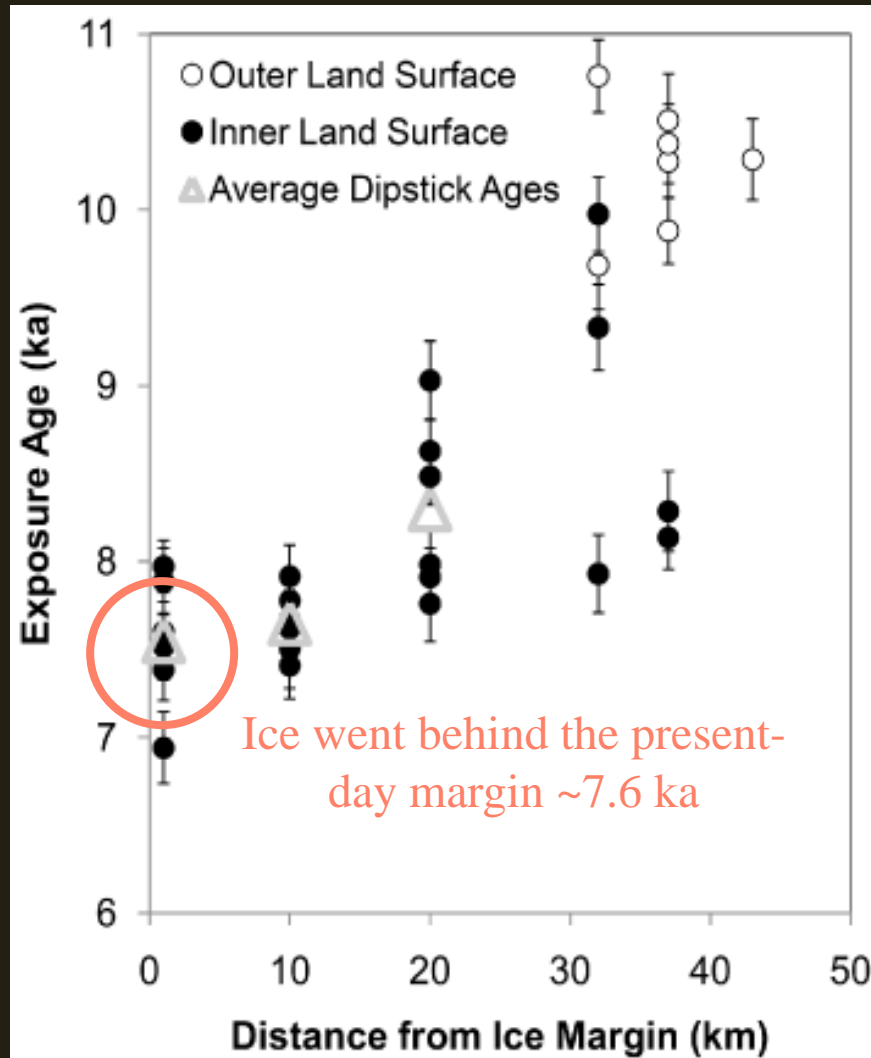
# Spatial Variability of Exposure Ages

There is no statistically significant difference between sample ages at high, medium, and low elevations.

The ice sheet thinned rapidly, at rates greater than what we can detect within the uncertainties associated with  $^{10}\text{Be}$  dating.

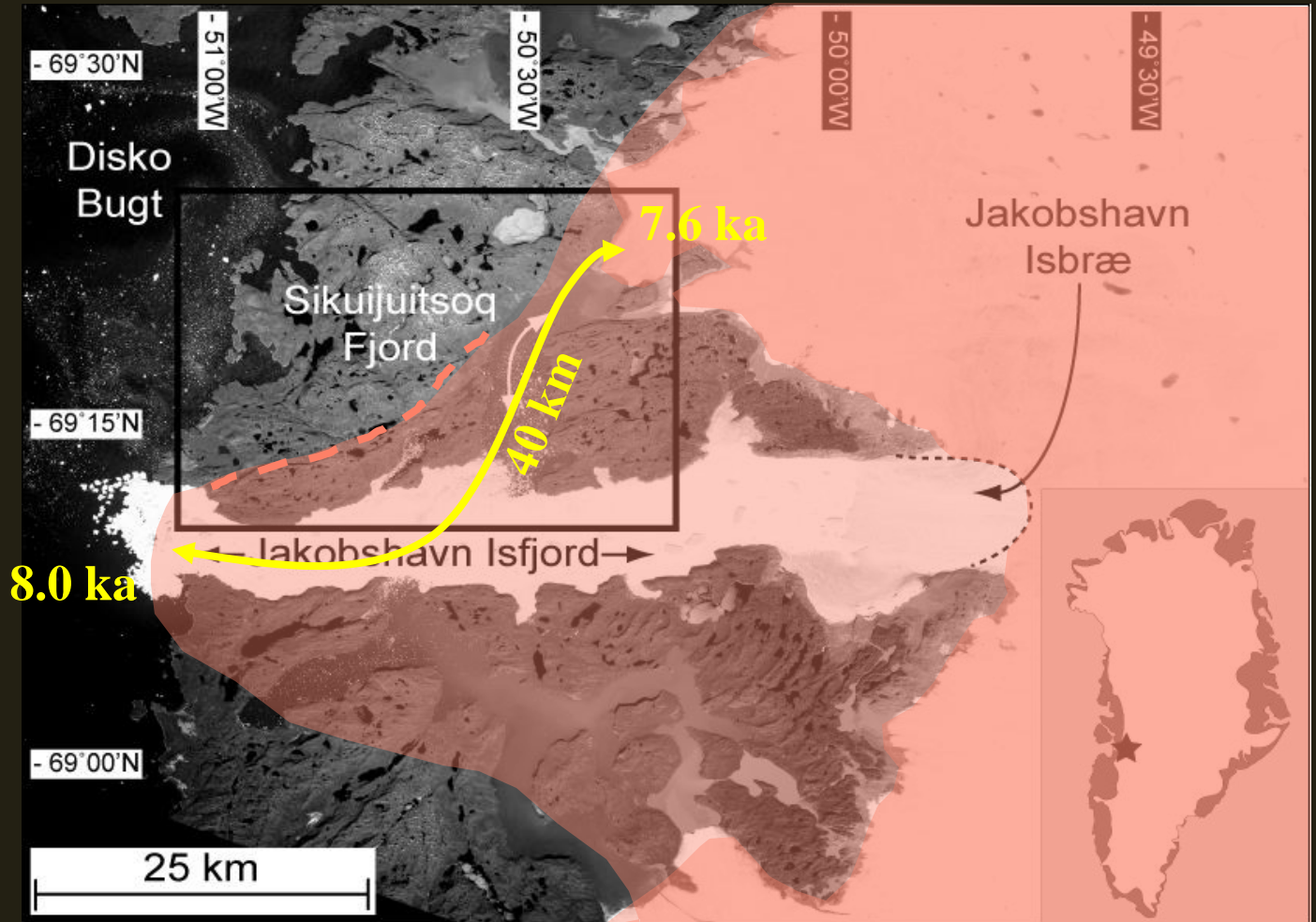


# Spatial Variability of Exposure Ages



How do we quantify ice margin retreat rates?

# Ice Margin Retreat



Retreat rate estimate:  
40 km of retreat in 400 yrs  $\approx$  100 m/yr



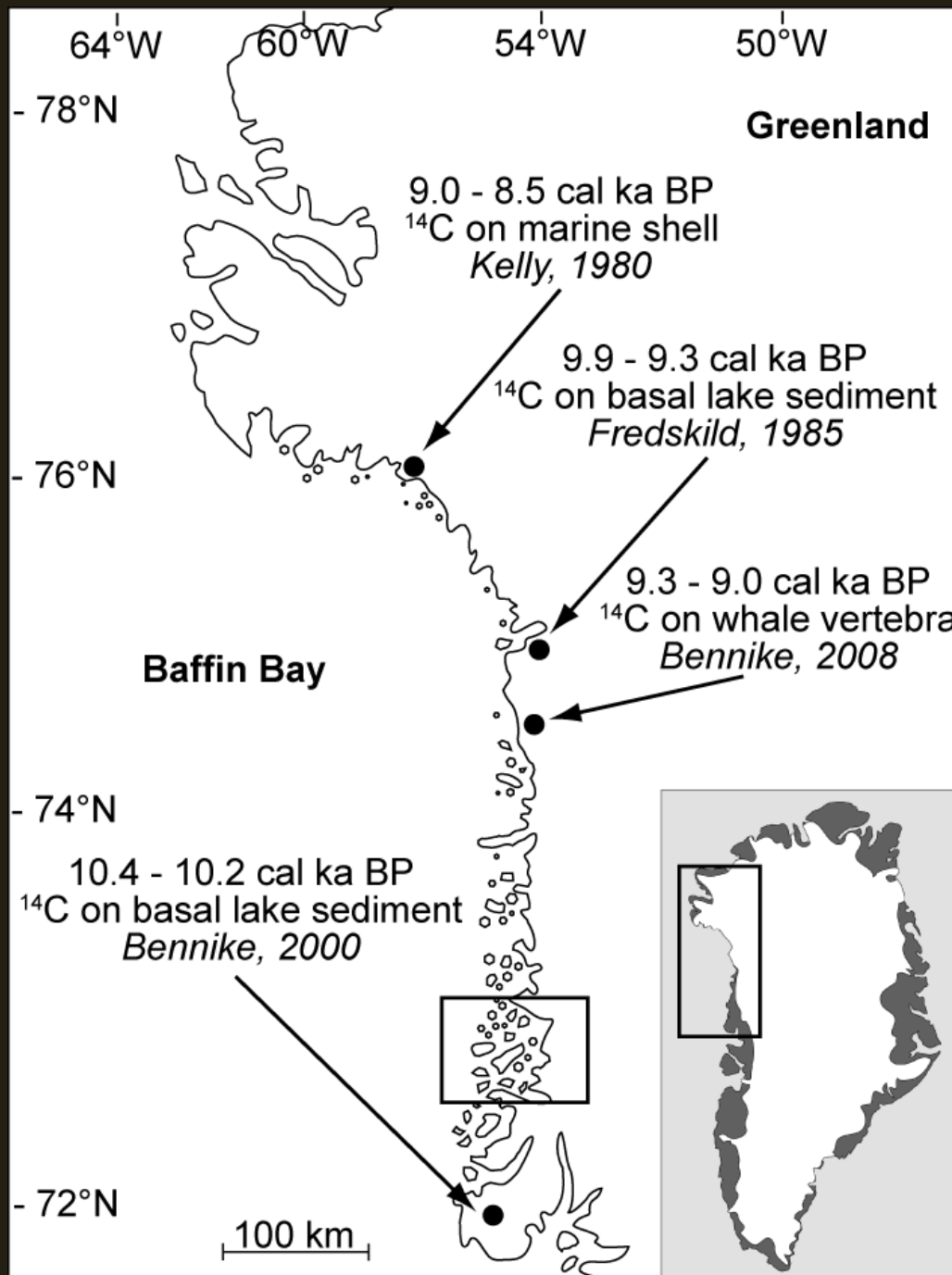
# Ilulissat Overview

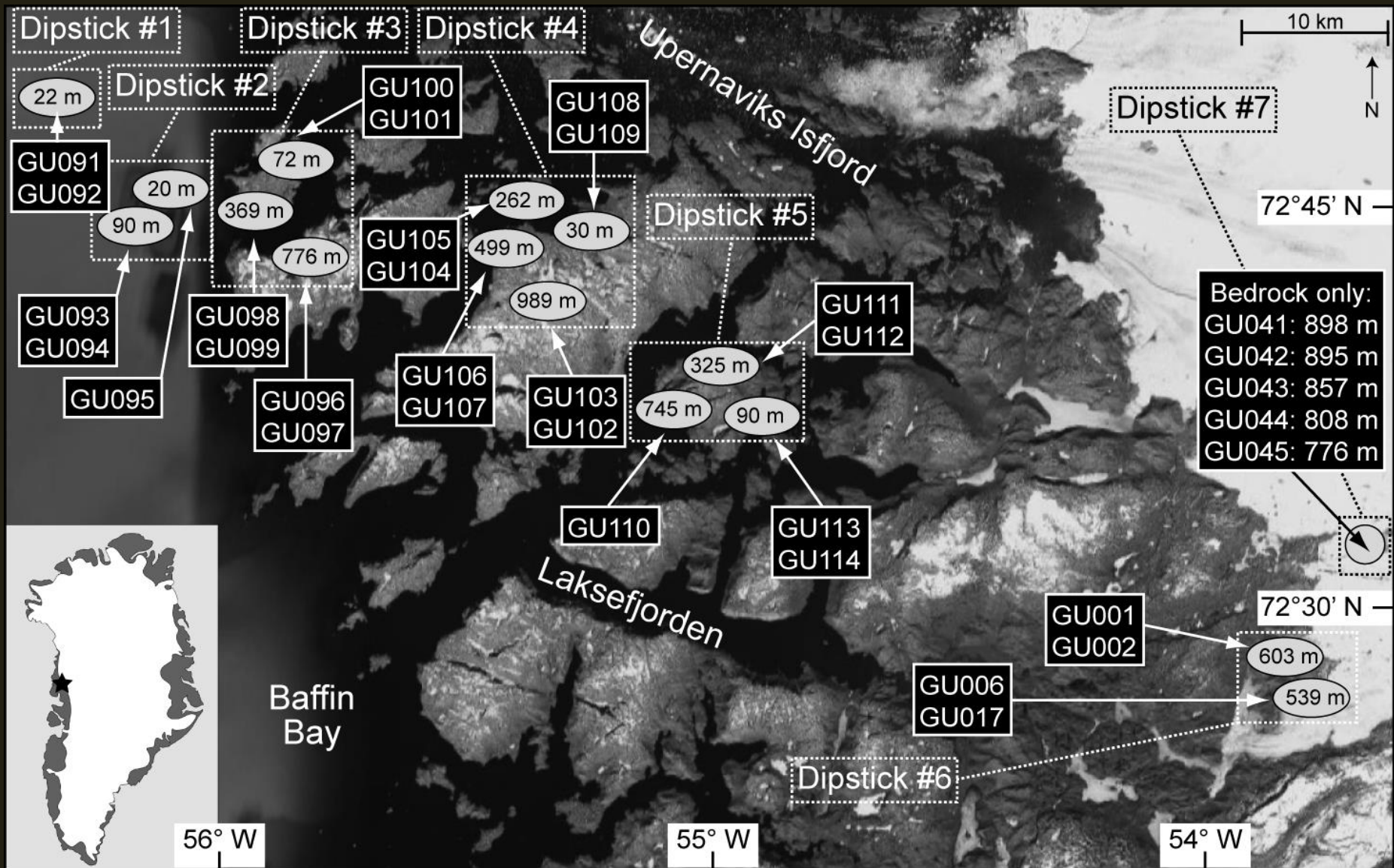
- All ages are Holocene
- Agreement between bedrock and boulder samples indicates little to no inheritance
- Efficient subglacial erosion ( $>2$  m)
- Deglaciation chronology is complicated due to the presence of a moraine: two land surfaces
- Fjord Stade moraine may be associated with the “8200 Event”
- Ice retreat began from the coast  $\sim 10.3$  ka, ice went behind the present-day margin  $\sim 7.6$  ka
- Ice retreated at  $\sim 100$  m/yr



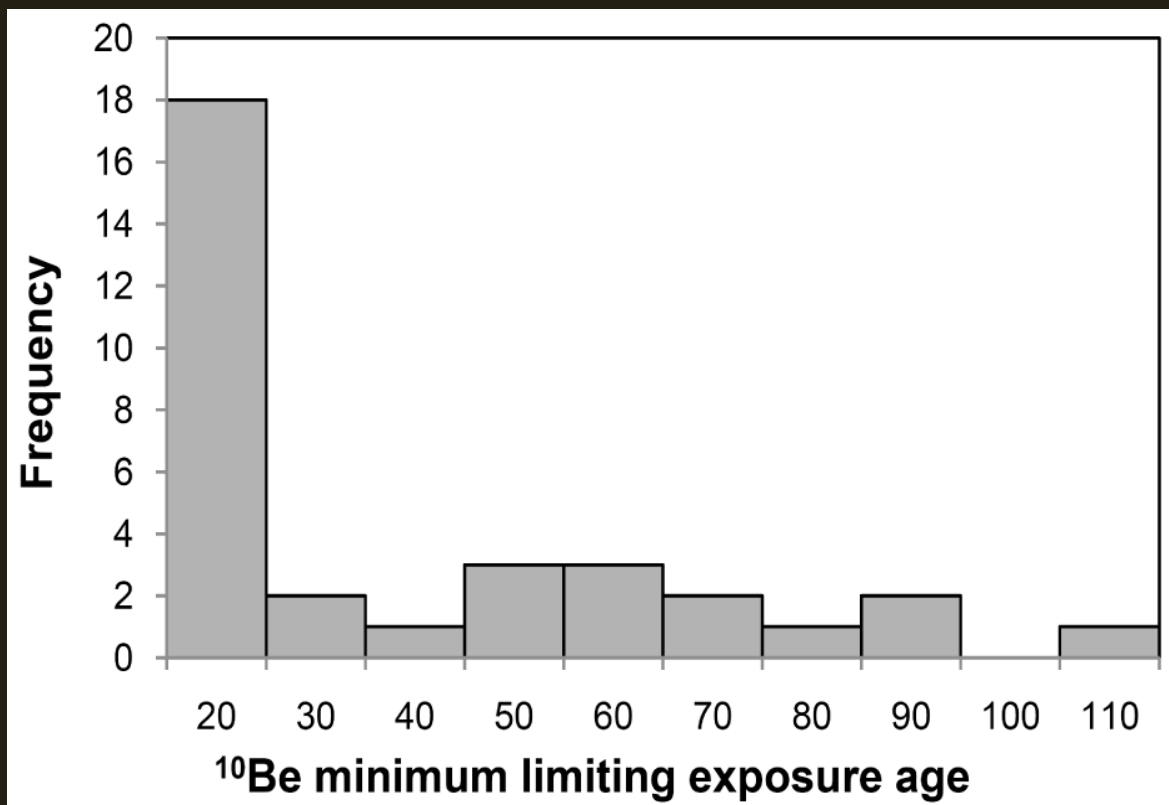


Thesis Part III:  
Upernavik, Greenland



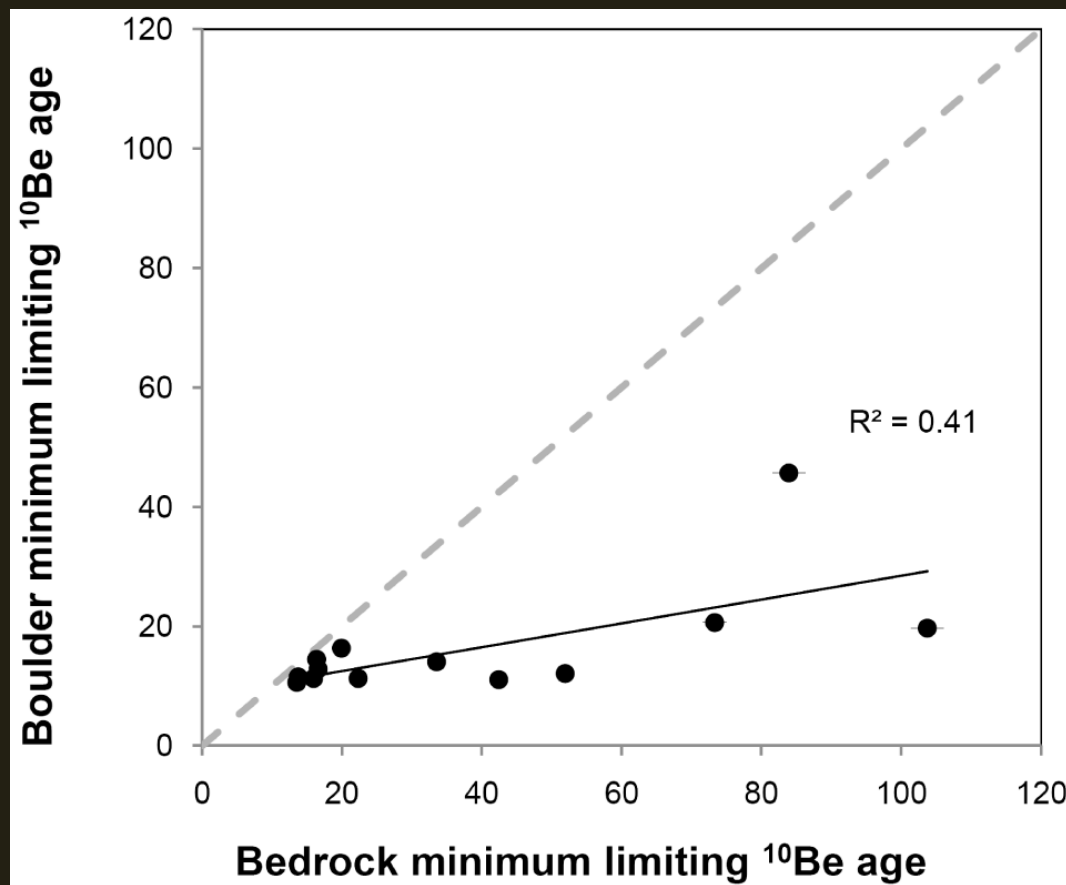


# Sample Ages



Something is going on here... why do we have such old exposure ages?

# Bedrock/Boulder Comparison



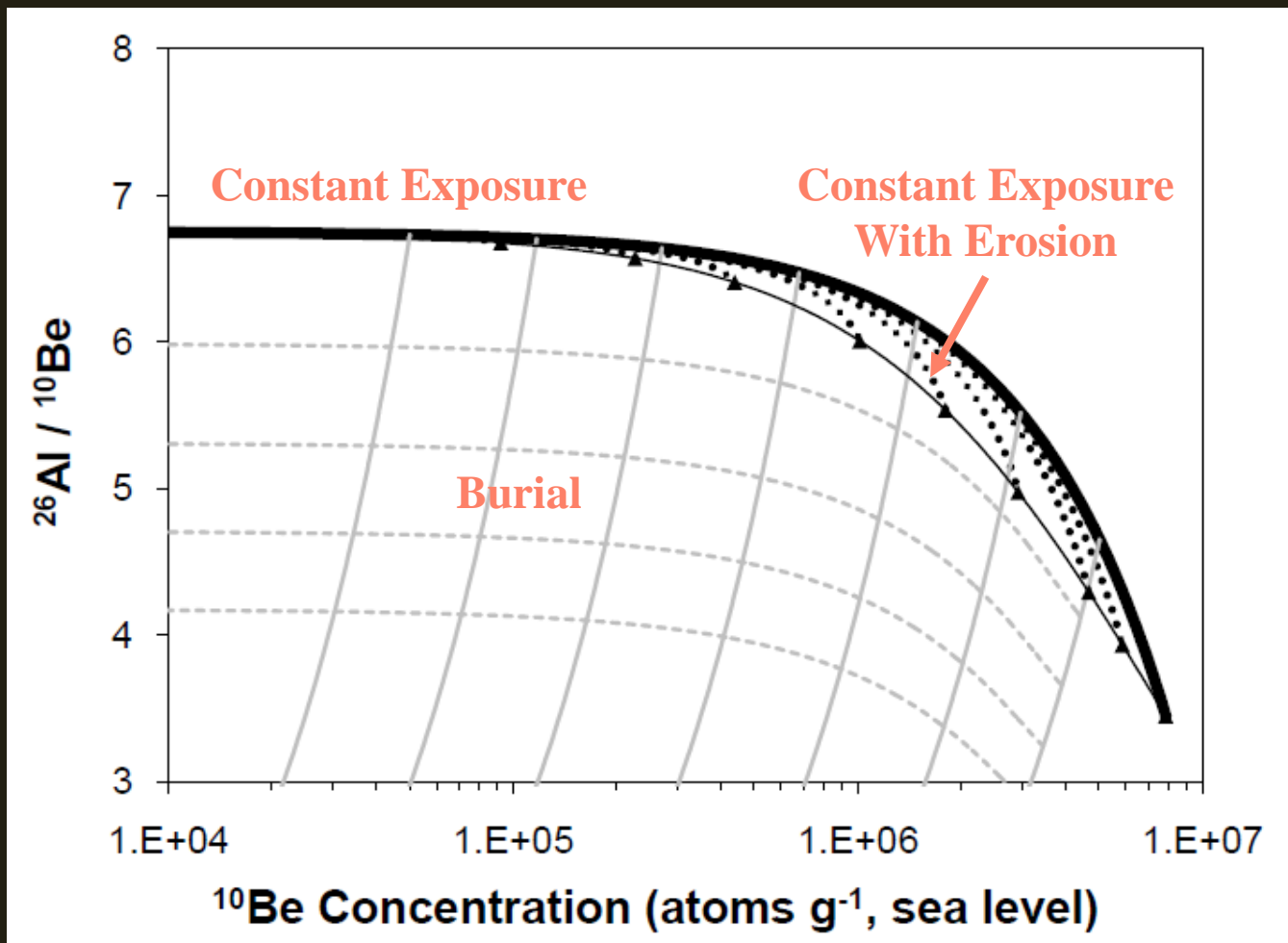
Bedrock samples are much older than paired boulder samples; inheritance is present. Glacial ice is non-erosive or weakly erosive.

# Using Cosmogenic $^{26}\text{Al}$

- Higher production rate than  $^{10}\text{Be}$
- Shorter half-life than  $^{10}\text{Be}$
- The two isotopes behave differently when burial occurs and production ceases

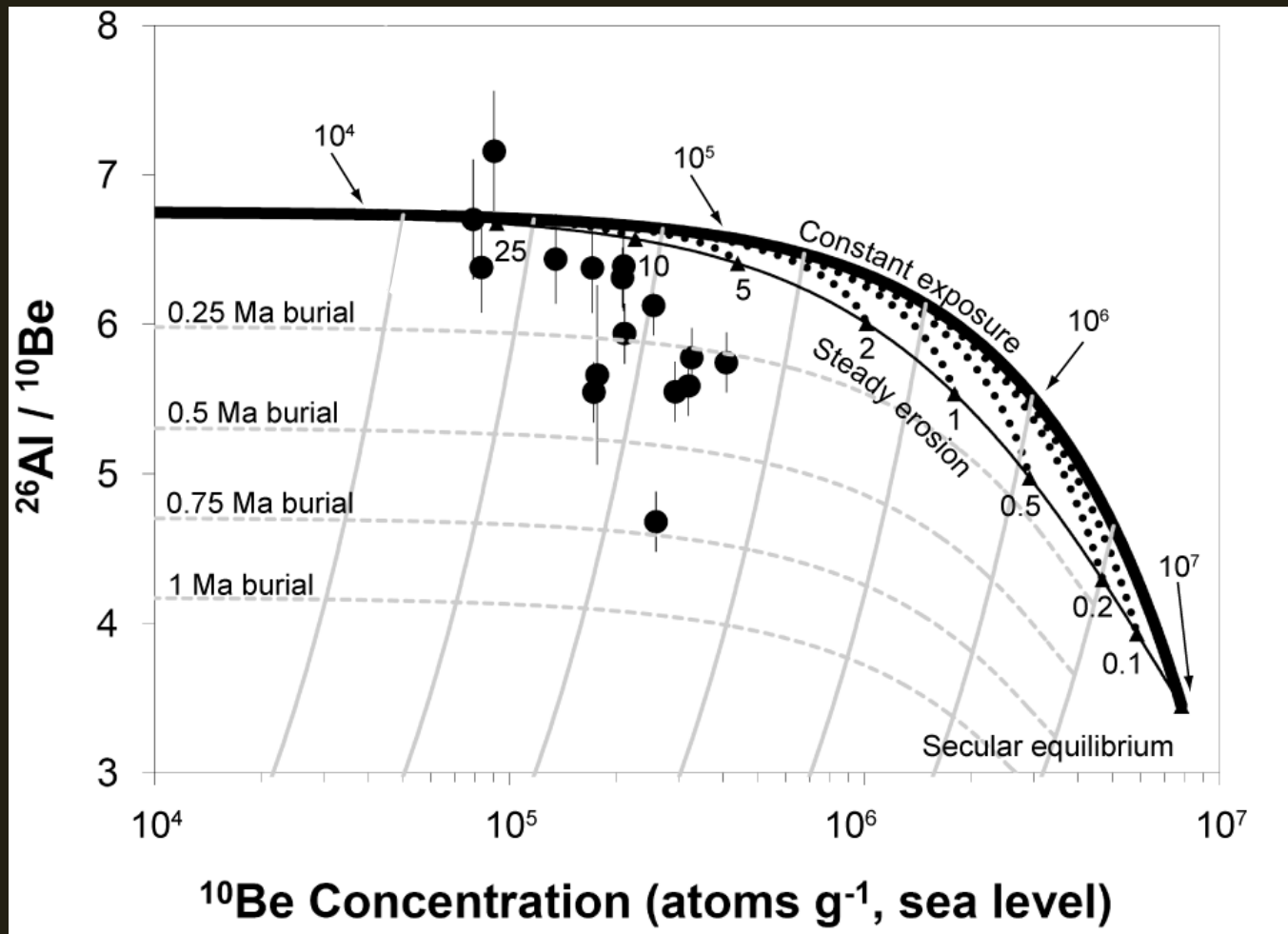


# The Two-Isotope Plot





# The Two-Isotope Plot



Samples have experienced both exposure and burial. They have long total histories.

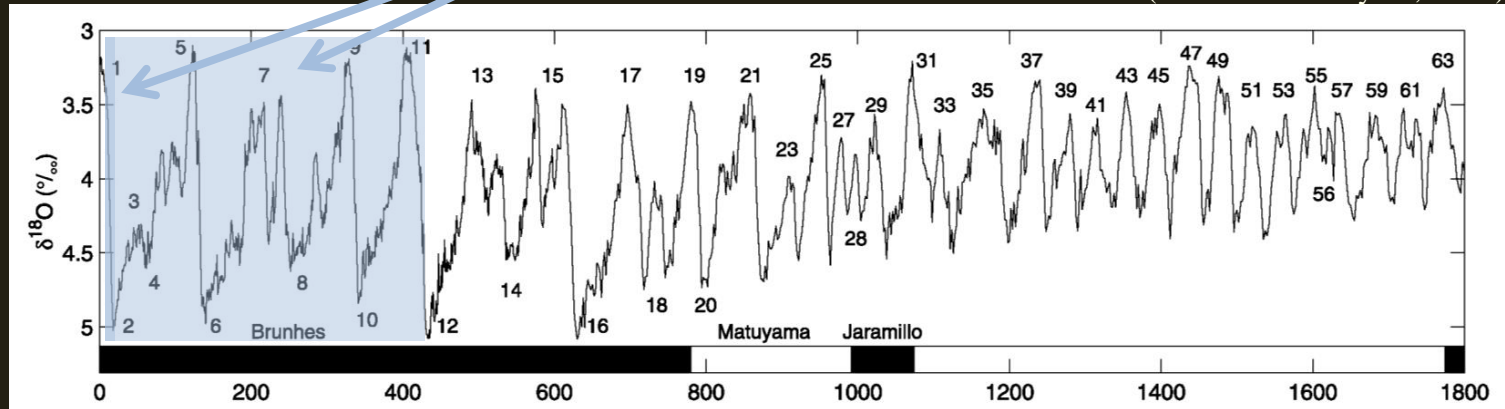
# Landscape History

Landscape history represented  
by Upernavik samples

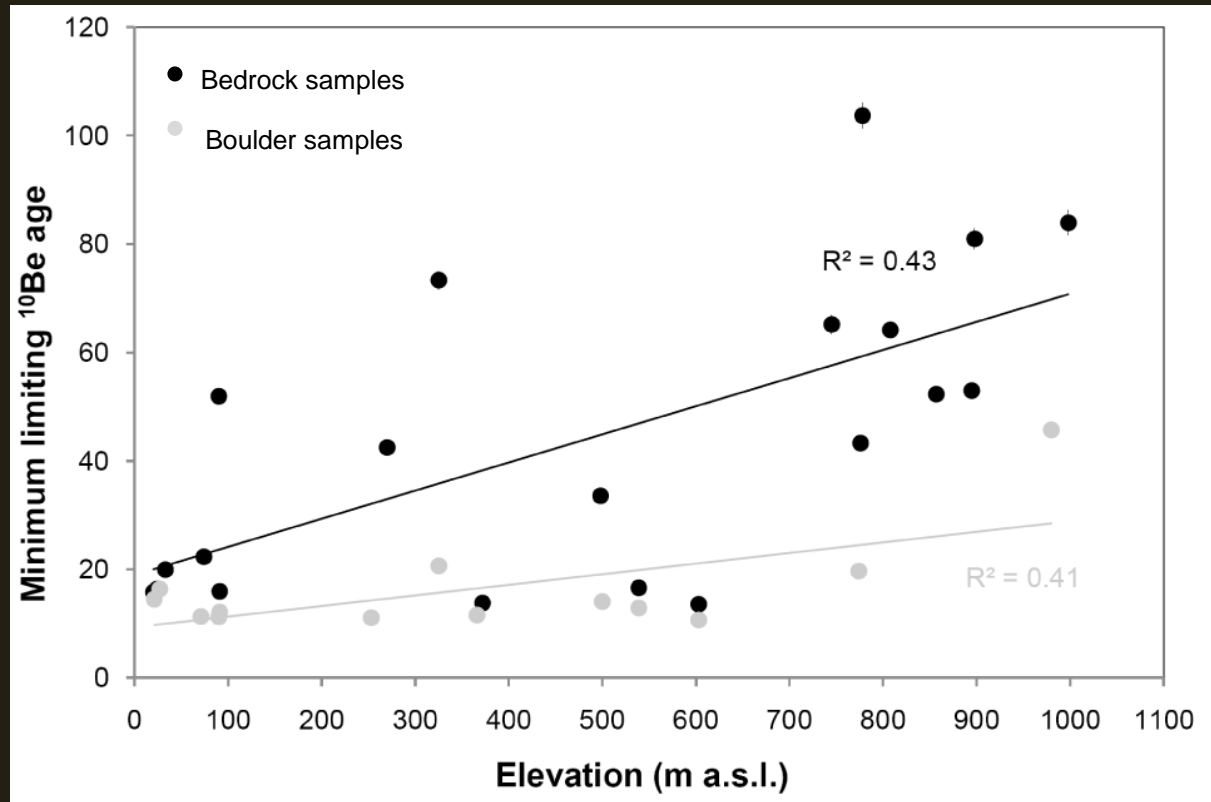
Warmer



Colder



# Inheritance and Elevation



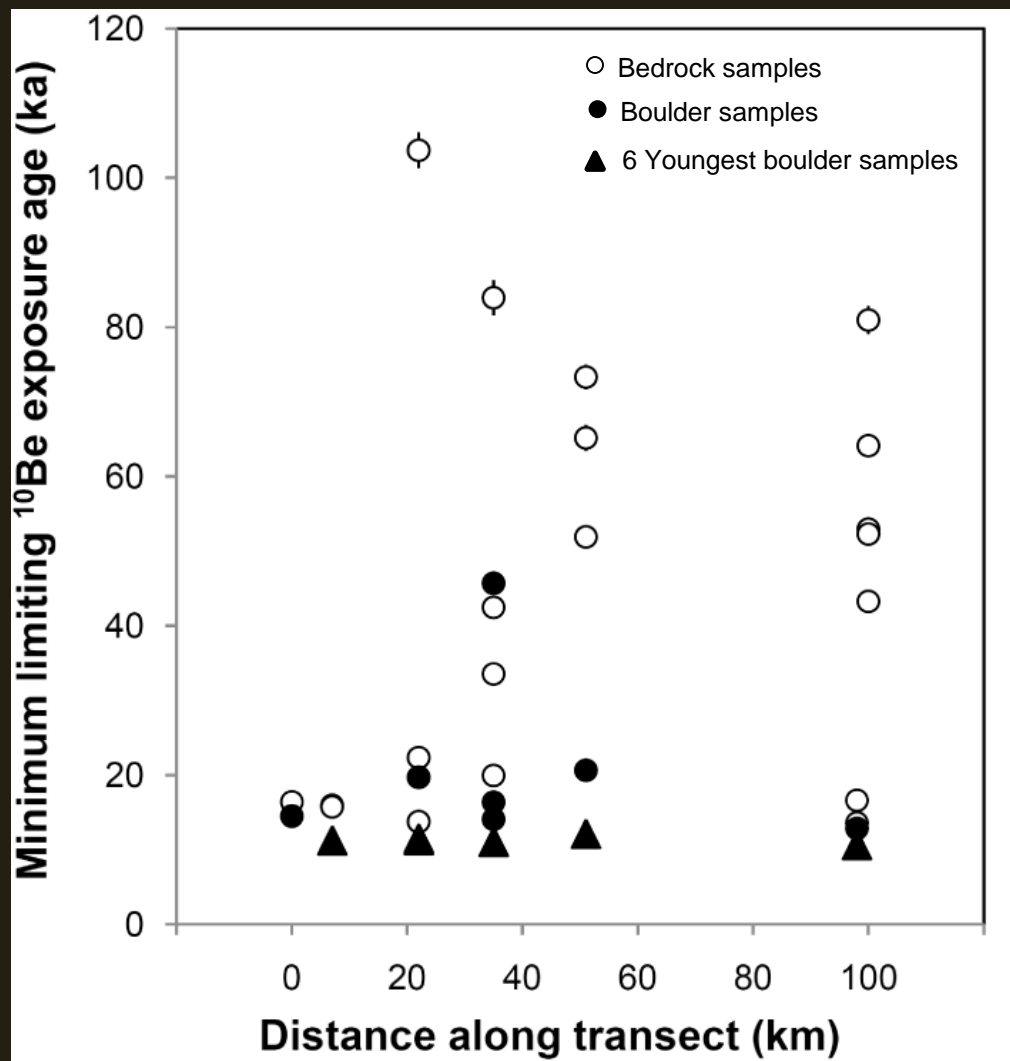
There is more inheritance at higher elevations.  
High-elevation ice is less erosive.

# Subglacial Erosion and Elevation

- There must be meltwater in order for ice to perform erosion
- “Warm-based” ice has meltwater at its bed due to warmer basal temperatures
  - With thicker ice, overlying weight decreases the pressure melting point
  - Low elevations
- “Cold-based” ice has no meltwater at its bed due to colder basal temperatures
  - With thinner ice, overlying weight is not sufficient to decrease the pressure melting point
  - High elevations

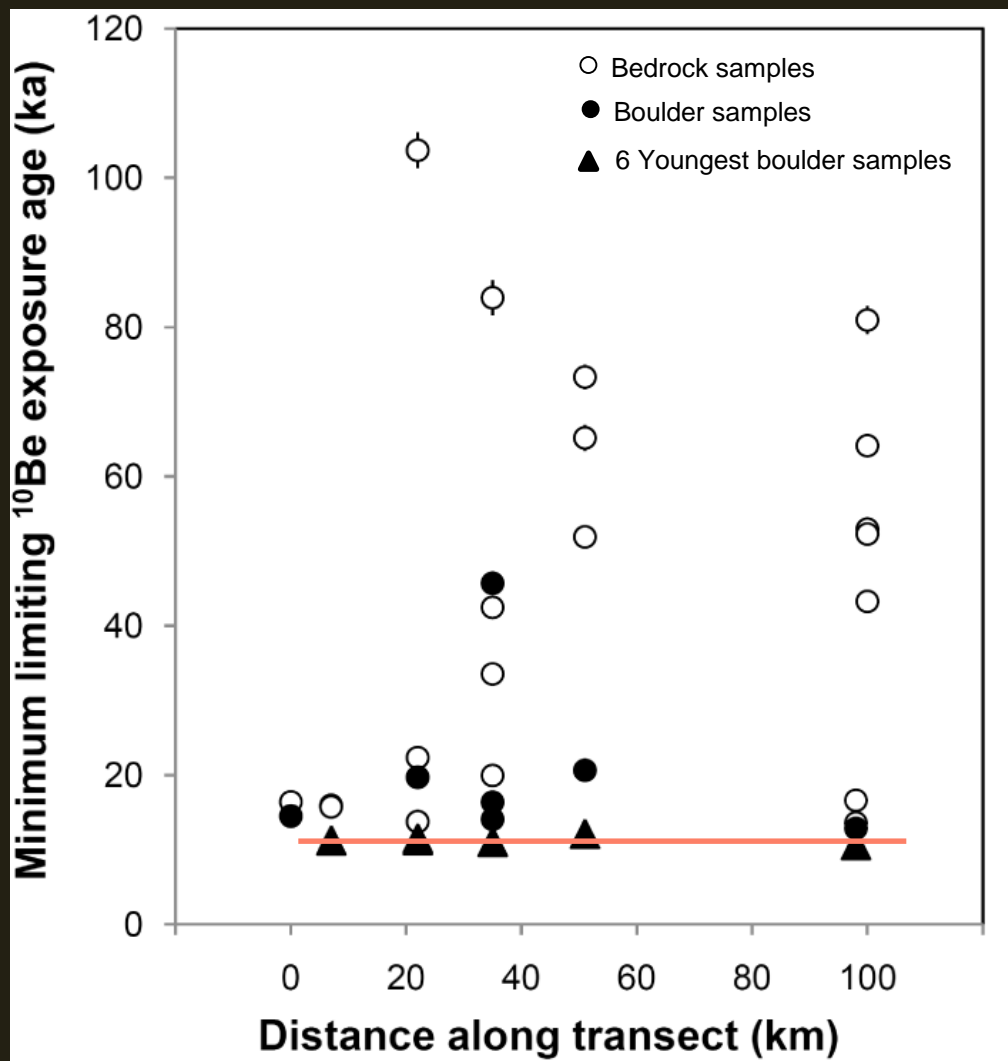


# Spatial Variability of Exposure Ages



How do we quantify ice margin retreat rates?

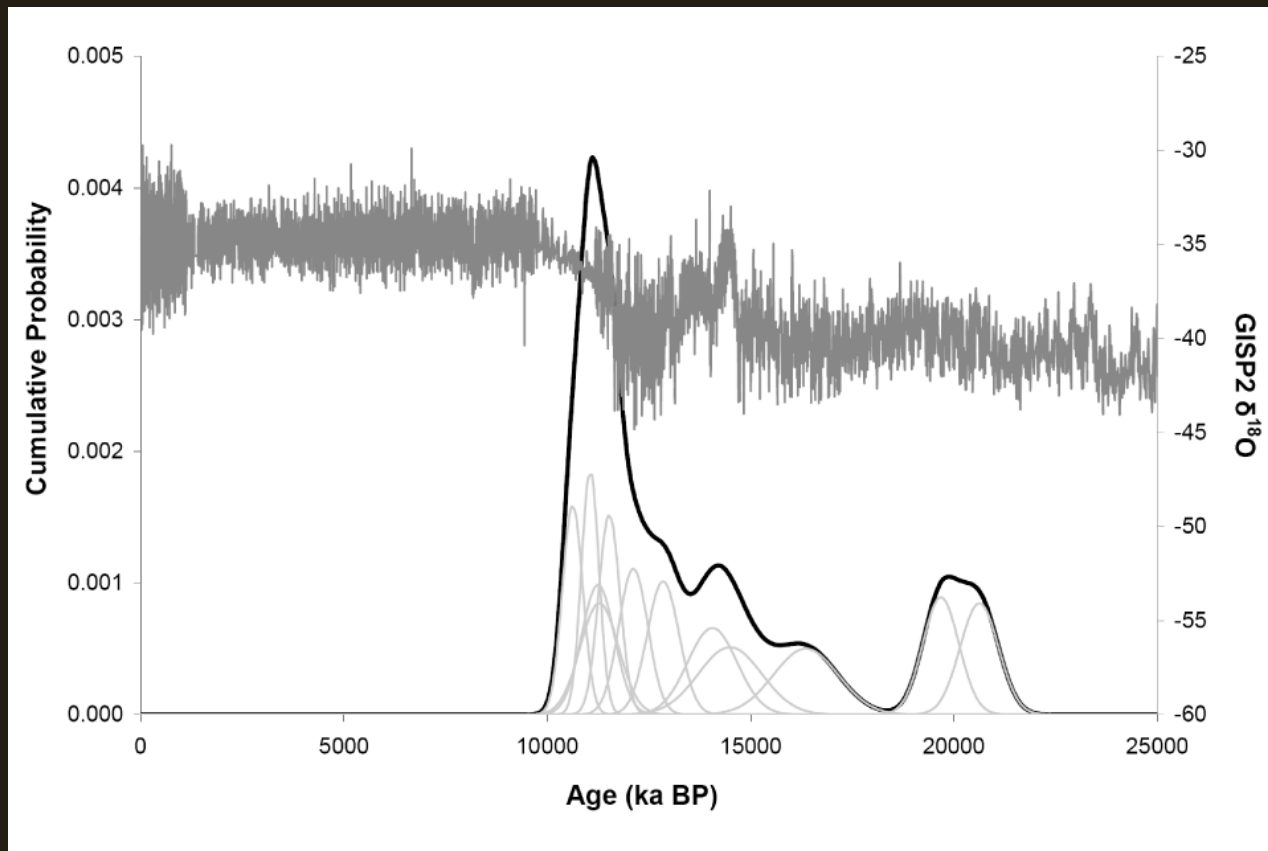
# Spatial Variability of Exposure Ages



6 youngest samples have statistically indistinguishable ages of  $11.3 \pm 0.5$  ka

Numerical modeling yields a statistically most likely retreat rate of  $\sim 170$  m/yr

# Paleoclimate Context



Rapid loss of ice may have occurred during warming after the Younger Dryas

# Upernavik Overview

- Most ages are old, few are Holocene
- Poor agreement between bedrock and boulder samples indicates ample inheritance
- Low subglacial erosion rates
- Ice retreat occurred rapidly at  $\sim 11.3 \pm 0.5$  ka
- Ice retreated at  $\sim 170$  m/yr
- Rapid ice retreat may have coincided with warming after the Younger Dryas







# Conclusions & The Big Picture



# Comparisons Between Sites

<b>ILULISSAT</b>	<b>UPERNAVIK</b>

High degree of variability between two sites  
only 500 km apart

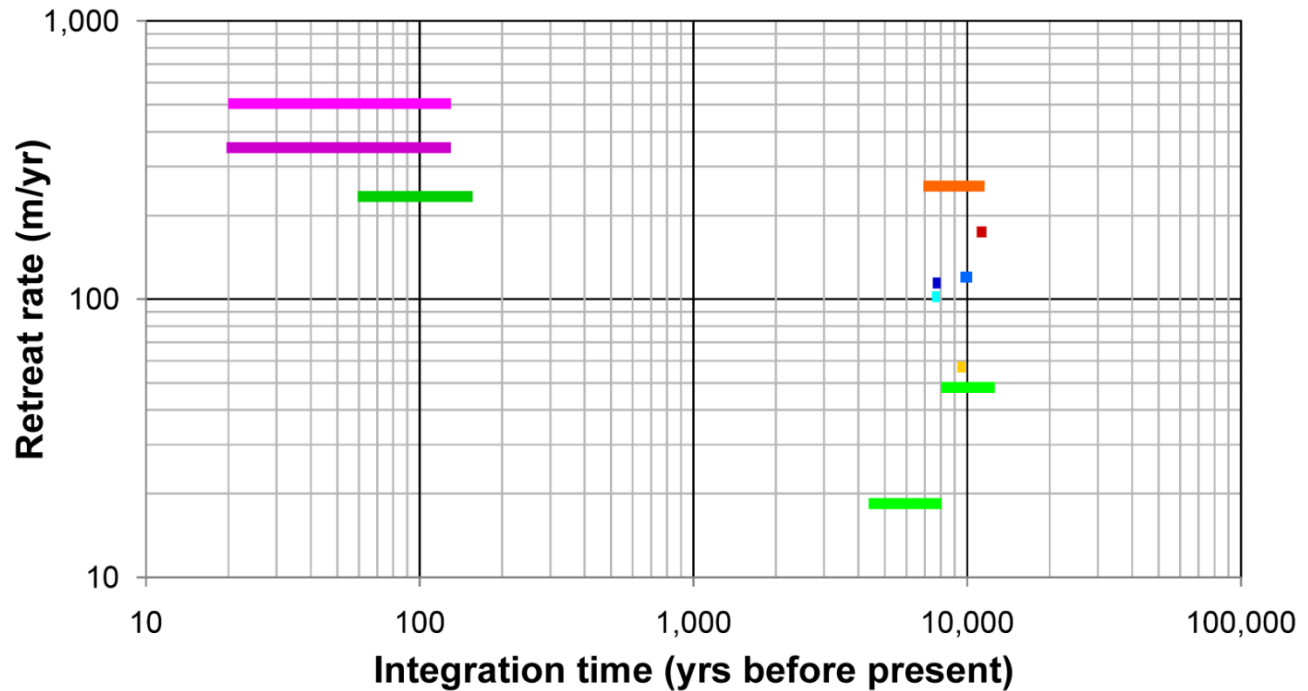
# What Does It All Mean?

## *Subglacial Erosion Rates*

- Subglacial erosion efficiency is controlled by local or regional factors
- Ice can be incredibly erosive, leading to sculpted, fresh landscapes
- Ice can be non-erosive, leading to old, heavily-weathered landscapes



# Ice Retreat Rate Comparisons



— = Sikuijuitsoq Fjord, West Greenland  
(Corbett et al., 2011)

— = Disko Bugt, West Greenland  
(Long and Roberts, 2003)

— = Jakobshavn Isfjord, West Greenland  
(Young et al., in press)

— = Jakobshavn Isfjord, West Greenland  
(Csatho et al., 2008)

— = Sisimiut Fjord, West Greenland  
(Rinterknecht et al., 2008)

— = McCarthy Glacier, AK, USA  
(Wiles and Calkin, 1993)

— = Icy Bay Glacier System, AK, USA  
(Porter, 1989)

— = Sam Ford Fjord, Baffin, Canada  
(Briner et al., 2009)

— = Laurentide Ice Sheet, USA  
(Andrews et al., 1973)

— = Upernavik, West Greenland  
(Corbett et al., *In Preparation*)



# What Does It All Mean?

## *Ice Retreat*

- Ice retreat rates are controlled by local or regional factors
- When the ice margin is constrained, retreat rates are limited; when the ice margin is unconstrained, retreat rates are more rapid
- Retreat rates of floating margins can be an order of magnitude faster than retreat rates of grounded margins
- Ice-loss through calving has important implications for future sea-level rise



# What Does It All Mean?

## *Rates*

Many geologic processes take place over *geologic time scales*.

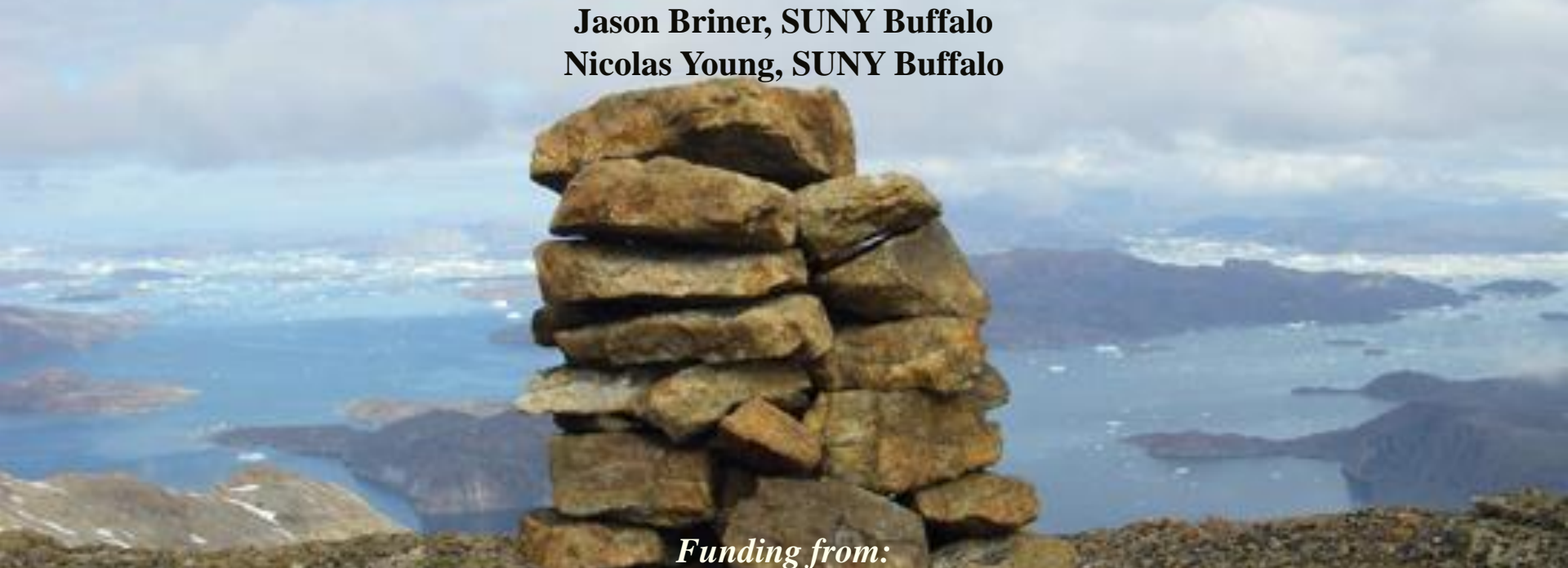
- Continents move at cm/yr
- Mountains erode at fractions of a mm/yr

Ice retreat in Greenland has, and will again, retreat over *human time scales*.

- Ice retreats at hundreds of m/yr



**Paul Bierman, University of Vermont**  
**Tom Neumann, NASA**  
**Joseph Graly, University of Wyoming**  
**Stephen Wright, University of Vermont**  
**Shelly Rayback, University of Vermont**  
**Bob Finkel, Lawrence Livermore**  
**Dylan Rood, Lawrence Livermore**  
**Jason Briner, SUNY Buffalo**  
**Nicolas Young, SUNY Buffalo**



*Funding from:*  
**NSF (ARC-0713956)**  
**NSF Graduate Research Fellowship**  
**University of Vermont**  
*Field support from:*  
**CH2MHILL**  
*Laboratory support from:*  
**Lawrence Livermore National Laboratory**





**UVM Geology Faculty**  
**UVM Geology Grad Students**  
**Will Hackett**  
**Eric Portenga**  
**Luke Reusser**  
**Charles Trodick**



**Pam and Pat Corbett**  
**Middlebury Friends**  
**Middlebury Geology Faculty**





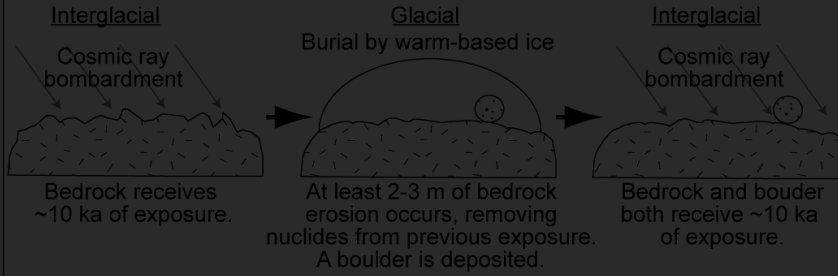




# Burial Scenarios

YOUNG BOULDER  
YOUNG BEDROCK

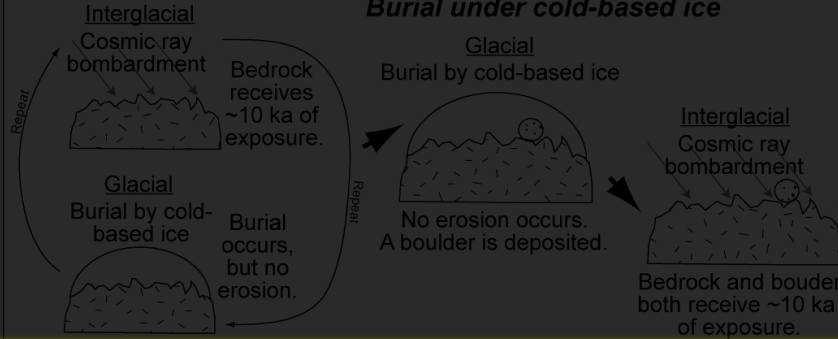
## Burial under warm-based ice



**Results**  
Boulder sample age: ~10 ka  
Bedrock sample age: ~10 ka  
Boulder  $^{26}\text{Al}/^{10}\text{Be}$  ratio: ~6.75  
Bedrock  $^{26}\text{Al}/^{10}\text{Be}$  ratio: ~6.75

YOUNG BOULDER  
OLD BEDROCK

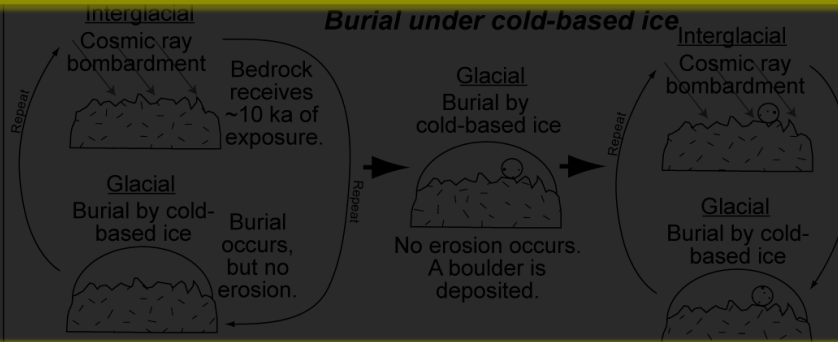
## Burial under cold-based ice



**Results**  
Boulder sample age: ~10 ka  
Bedrock sample age: 10's to 100's of ka  
Boulder  $^{26}\text{Al}/^{10}\text{Be}$  ratio: ~6.75  
Bedrock  $^{26}\text{Al}/^{10}\text{Be}$  ratio: <6.75

YOUNG BOULDER  
OLD BEDROCK

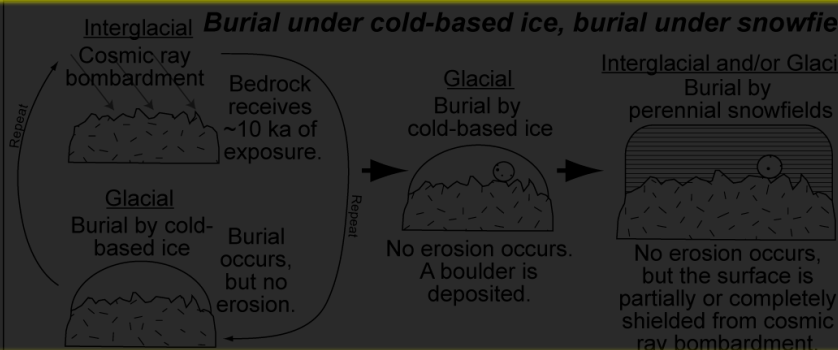
## Burial under cold-based ice



**Results**  
Boulder sample age: 10's of ka  
Bedrock sample age: 100's of ka  
Boulder  $^{26}\text{Al}/^{10}\text{Be}$  ratio: <6.75  
Bedrock  $^{26}\text{Al}/^{10}\text{Be}$  ratio: <<6.75

YOUNG BOULDER  
OLD BEDROCK

## Burial under cold-based ice, burial under snowfields



**Results**  
Boulder sample age: 10's of ka  
Bedrock sample age: 100's of ka  
Boulder  $^{26}\text{Al}/^{10}\text{Be}$  ratio: <6.75  
Bedrock  $^{26}\text{Al}/^{10}\text{Be}$  ratio: <<6.75