

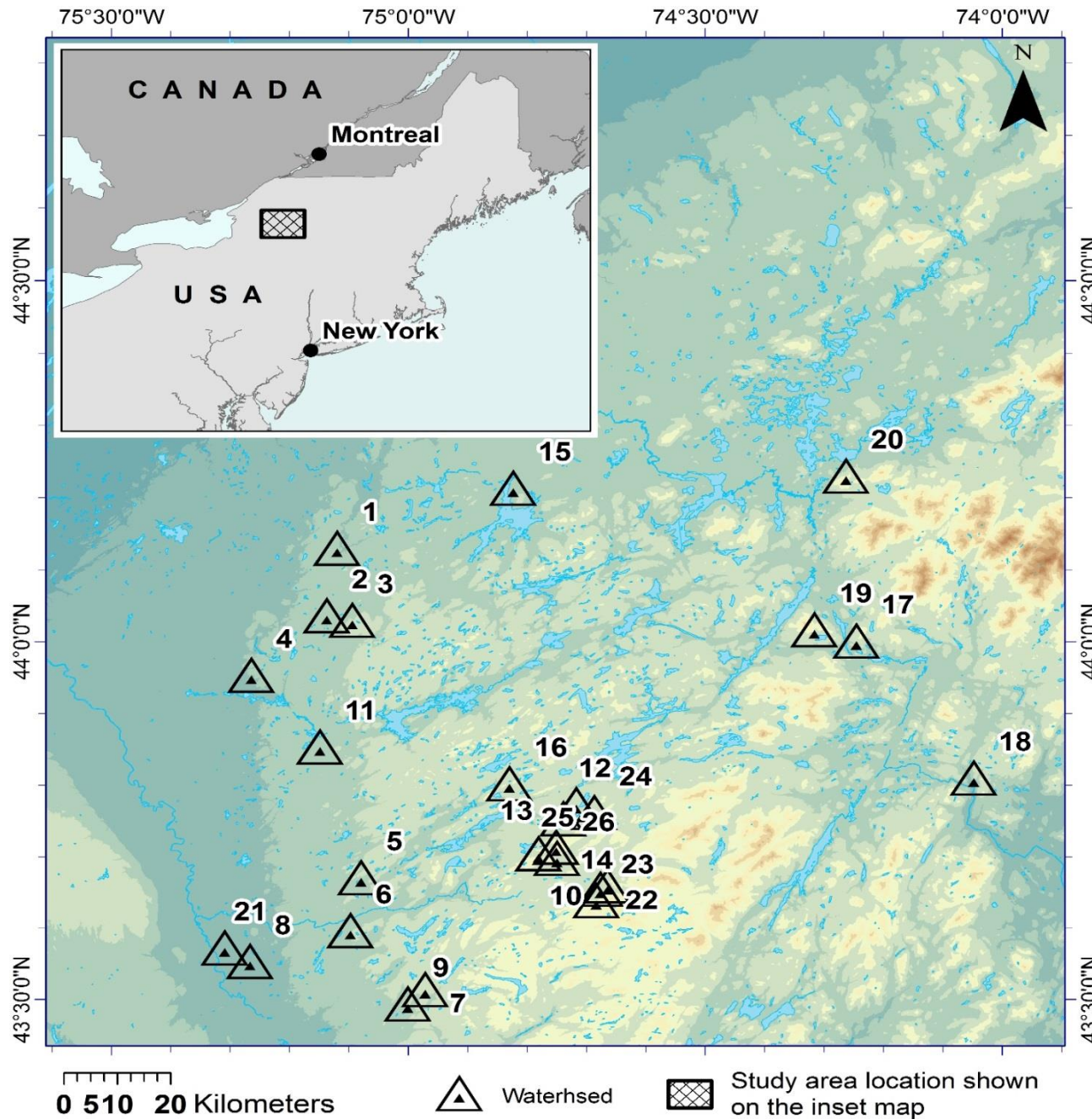
“ Trees sensitivity to Warming is controlled by Turnover Time of Forest Floor”

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Reasons for this work

- 1. There are only a few (mostly agricultural) studies about linkage between spring phenology and soil-related variables. Here we had an opportunity to test connection between spring phenology of natural growths with edaphic factors.
- 2. There is no any studies on potential linkage of plants sensitivity to warming with edaphic factors.



26 small, deciduous watersheds. Sampled: 2004, 2009

Exchangeable chemistry, C_{org} , N, Oa thickness, bulk density

Remote sensing data : Start of Season (**SoS**), End of Season (**EoS**), **NPP**. Source: Moderate Resolution Imaging Spectroradiometer MOD13Q1, MOD17A3.

Scope of the work

- 1. Investigated potential linkage between SoS (first greenup dates during 2001-2009 period) and soil-related variables.
- 2. Estimated plants sensitivity to warming (β) as regression coefficient of SoS to spring (April or March) temperature during 2001-2009 period [days of SoS advance/retreat per 1 °C]
- 3. Investigated potential linkage of statistically significant estimates of β and soil-related variables.

Climatological variables

- Mean Annual Temperature (**MAT**), Annual Precipitation (**P**), Daily Maximum Temperature, Mean Daily Temperature.
- “Biological Temperature” (**BT**=5°C) and the Start of the Thermal Growing Season (**STGS**).
- In addition to STGS we did estimate β_T (sensitivity of thermal season to changes in spring temperature): regression coefficient between STGS and spring temperature : days of STGS advance/reatreat per 1° C.

IMPORTANT

β and β_T have the same dimension: [days per 1 °C.]

Additional, soil-related variable: Turnover time of Soil Organic Matter (SOM) in forest floor

$$\bullet \tau[\text{yr}] = \frac{\text{Stock of Corg in forest floor layer} \left[\frac{\text{kg}}{\text{m}^2} \right]}{\text{Litterfall} \left[\frac{\text{kg}}{\text{m}^2 \text{yr}} \right]} \quad (1)$$

$$\bullet \quad \text{Stock of Corg in forest floor} = C_{org} \times h \times \rho \quad (2)$$

- Where h , - the thickness, and, ρ , - bulk density of forest floor.
- The litterfall was estimated as function of NPP:

$$\bullet \quad \text{Litterfall} = \alpha \times \text{NPP} \quad (3)$$

- Where α is the litterfall fraction of NPP which varies with forest type and age within the range from 20 to 60%. In old-growth, hardwood stands of Great Lake region, this parameter equals to about 40% (Tang and Bolstad, 2005).

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RESULT #1

SoS

- SoS correlates with *Al*, CEC, *pH* at $p < 0.05$ and with MAT and thickness of Oa layer (*h*) at $p < 0.01$.
- **Covariance:** *Al* with exchangeable chemistry is causal!
- MAT with the elevation above sea level and with thickness of Oa layer (*h*),
- MAT is controlled by the elevation above sea level (the slope is close to the lapse rate at 6.2 °C /km);
- MULTIPLE REGRESSION MODEL reveals **0.37** days per 1 [meq/100g] of *Al*; **5.1** days per 1 °C of MAT, and **1.7** days per 1 cm of forest floor.
- **Covariance of MAT and the thickness of Oa layer can be causal!**

Thermal diffusion model of Oa layer

Delay (phase shift) of transmittance of the thermal signal from near surface atmosphere into the soil can be described as (Van Wijk, 1963; Jury, Horton, 2004):

- $\Delta t = h/\omega d$ (5)

- where ω is the angular frequency of surface temperature; d is the damping depth; and h is soil depth in forest floor. The d term can be calculated as a function of ω and the coefficient of thermal diffusivity (K) (Jury, Horton, 2004):

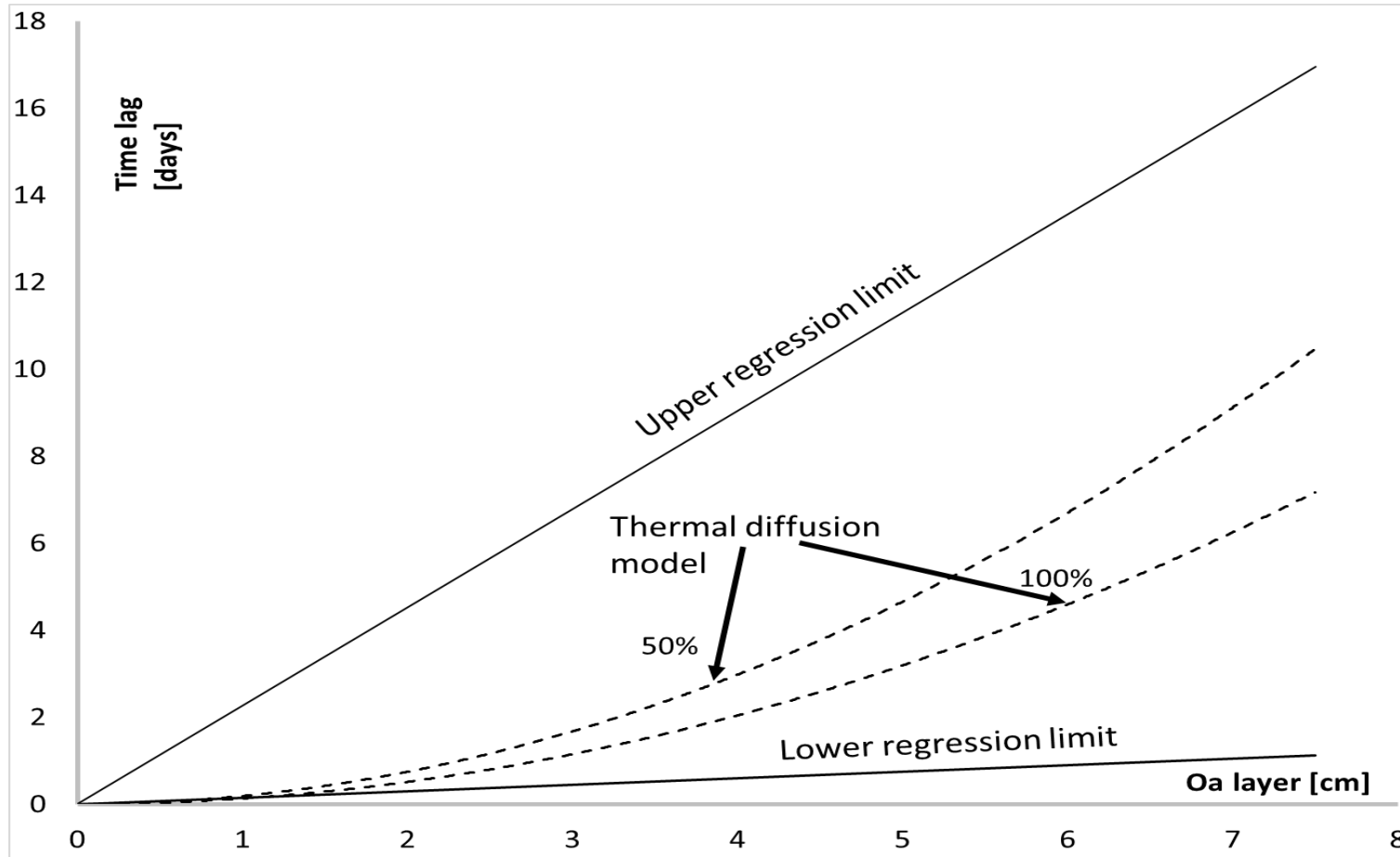
- $d = \sqrt{\frac{2K}{\omega}} = \sqrt{\frac{K\tau}{\pi}}$ (6)

- where τ is the time period (1 year). K values of the forest floor depend on bulk density and the proportion of organic matter relative to water and air.

Calculations of thermal diffusion coefficient

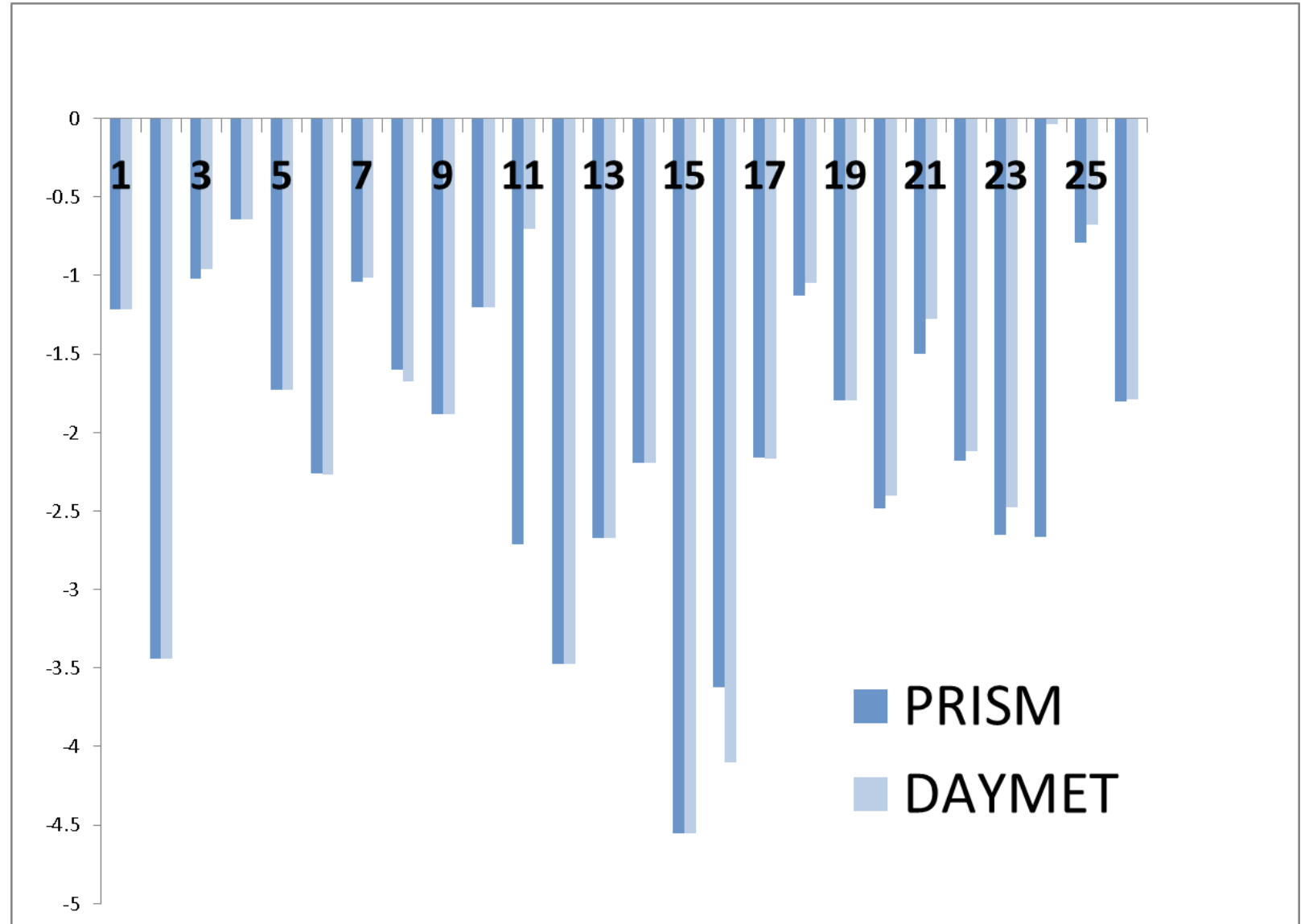
- Derived K values are about 5-fold smaller than the typical thermal diffusivity of snow (Oldroyd et al., 2013).
- According to our calculations, insulating capacity of the 10 cm of forest floor equals approximately to 0.5 meter of snow depth.

Experimental (95% confidence level) regression slopes and the thermal diffusion model results (100 and 50% saturation) of the delay of SoS relative to the thickness of forest floor (h).



Result #2

β



RESULT#3, β

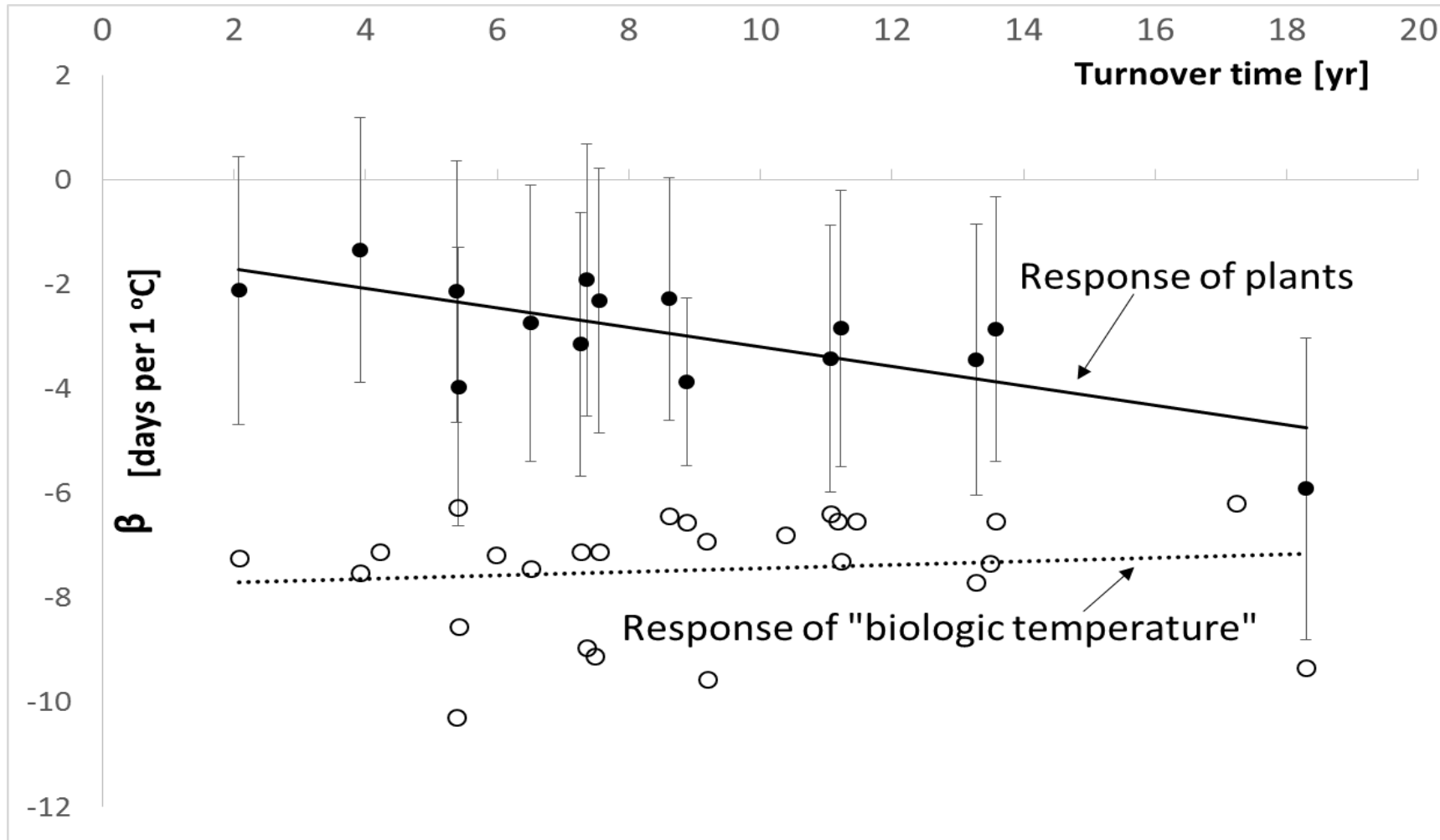
- We did build 3 multiple correlation models of **β and soil-related variables** :
- **1.** PRISM mean daily temperature data (τ , EoS); $r > 0.5$, $p < 0.05$, $n = 17$
- **2.** PRISM maximum temperature data (τ); $r > 0.8$; $p < 0.01$, $n = 17$
- **3.** DAYMET maximum temperature data (τ , MAT); $r > 0.7$, $p < 0.01$, $n = 17$

B increases with turnover time at the rate of about 0.1 days per 1°C per each 1 yr of turnover time.

(The whole range of turnover time estimates is from 3 to 20 years. The average of 9 years is very close to other independent estimates of turnover time in maple and oak hardwood stands).

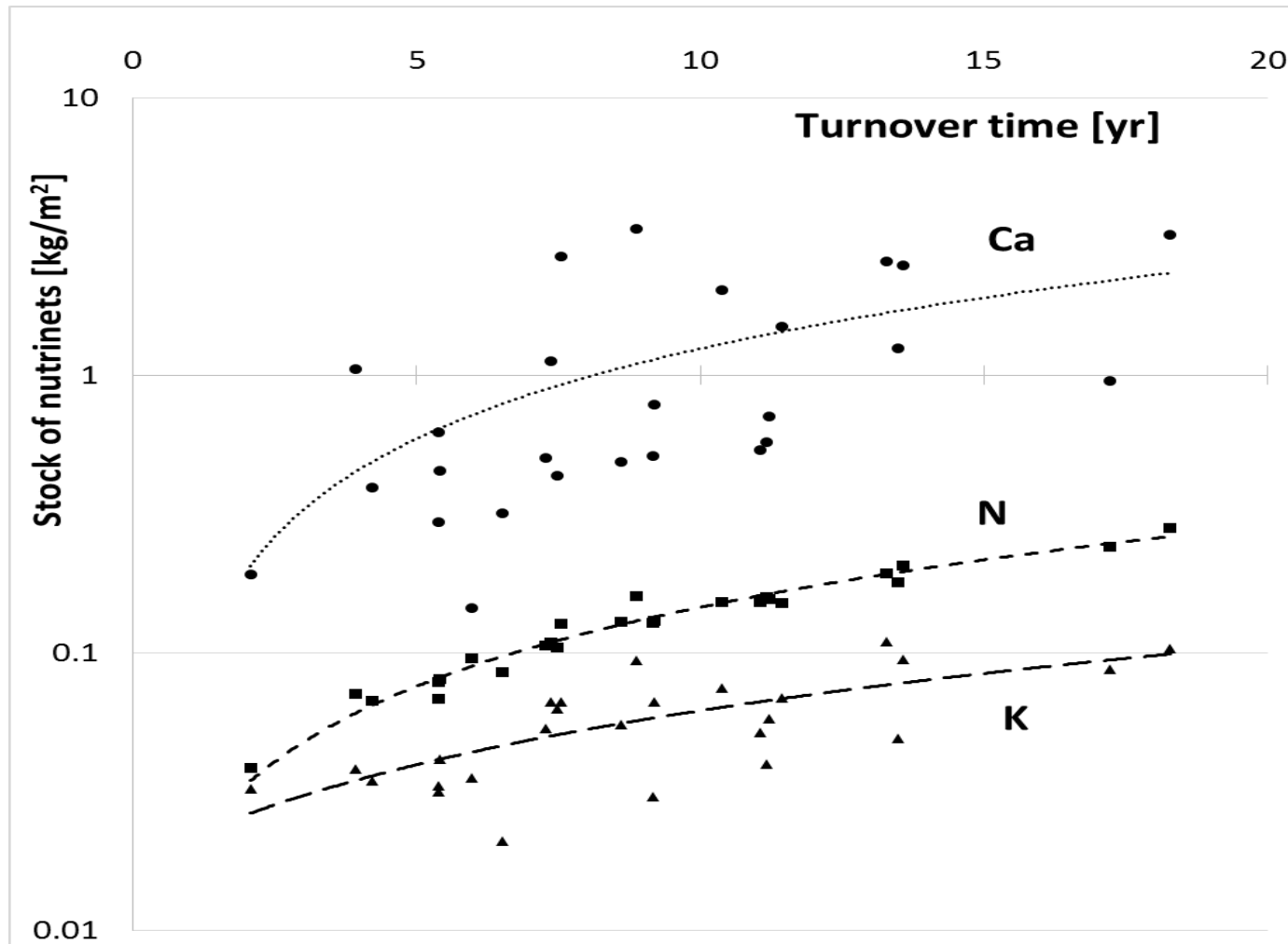
RESULT#3,

β



HYPOTHESIS 1: Soil Nutrients Limit Plants Response to Warming

HYPOTHEHSIS 2: Thick forest floor (slow turnover time) protects fine roots from late winter-early spring frost .



Conclusions

- 1. Spring phenology of similar hardwood stands in Adirondacks, in the region with relatively uniform climatic conditions, is controlled by soil acidification (A^{3+}) as well as by MAT and/or by the thickness of forest floor which acts as a very good thermal insulator, thus, delaying the spring thermal signal propagation into the soil.
- 2. Plants sensitivity to warming in Adirondacks is typical for deciduous forest ecosystem: 1 to 5 days advance of SoS per 1°C of spring warming.
- 3. The dominant factor controlling plants sensitivity to warming is turnover of forest floor. Most likely forest floor acts here as source of extra nutrients to supply extended growing season and/or as the last defense from late winter-early spring frost as snow thickness declines with the warming.